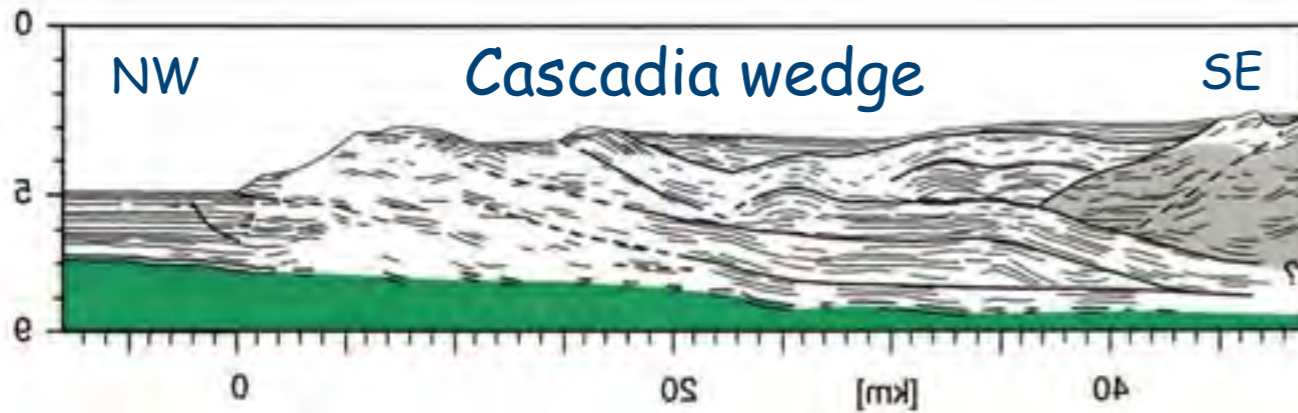


11- Kollision zonen

Akkretion

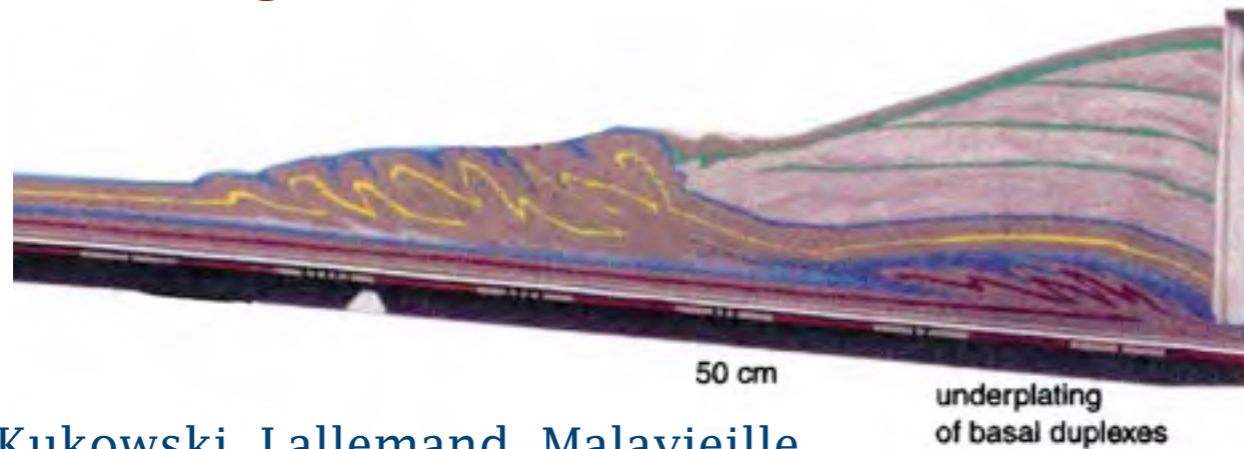


Mechanisms of accretion



Gutscher, M.A., Kukowski, N., Malavieille, J., and Lallemand, S., 1998, JGR

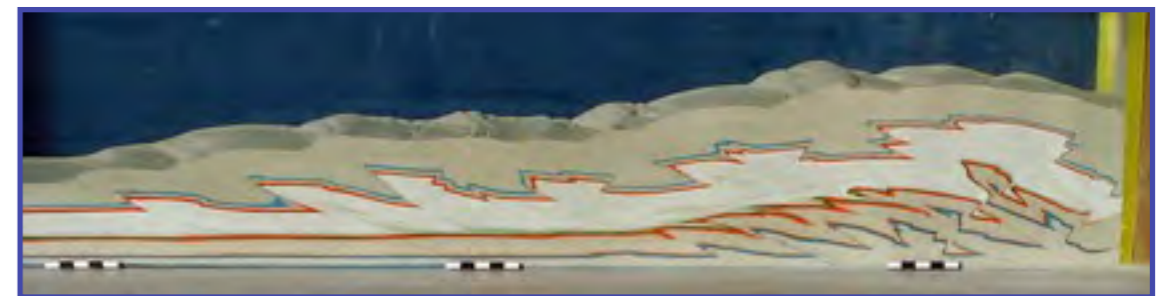
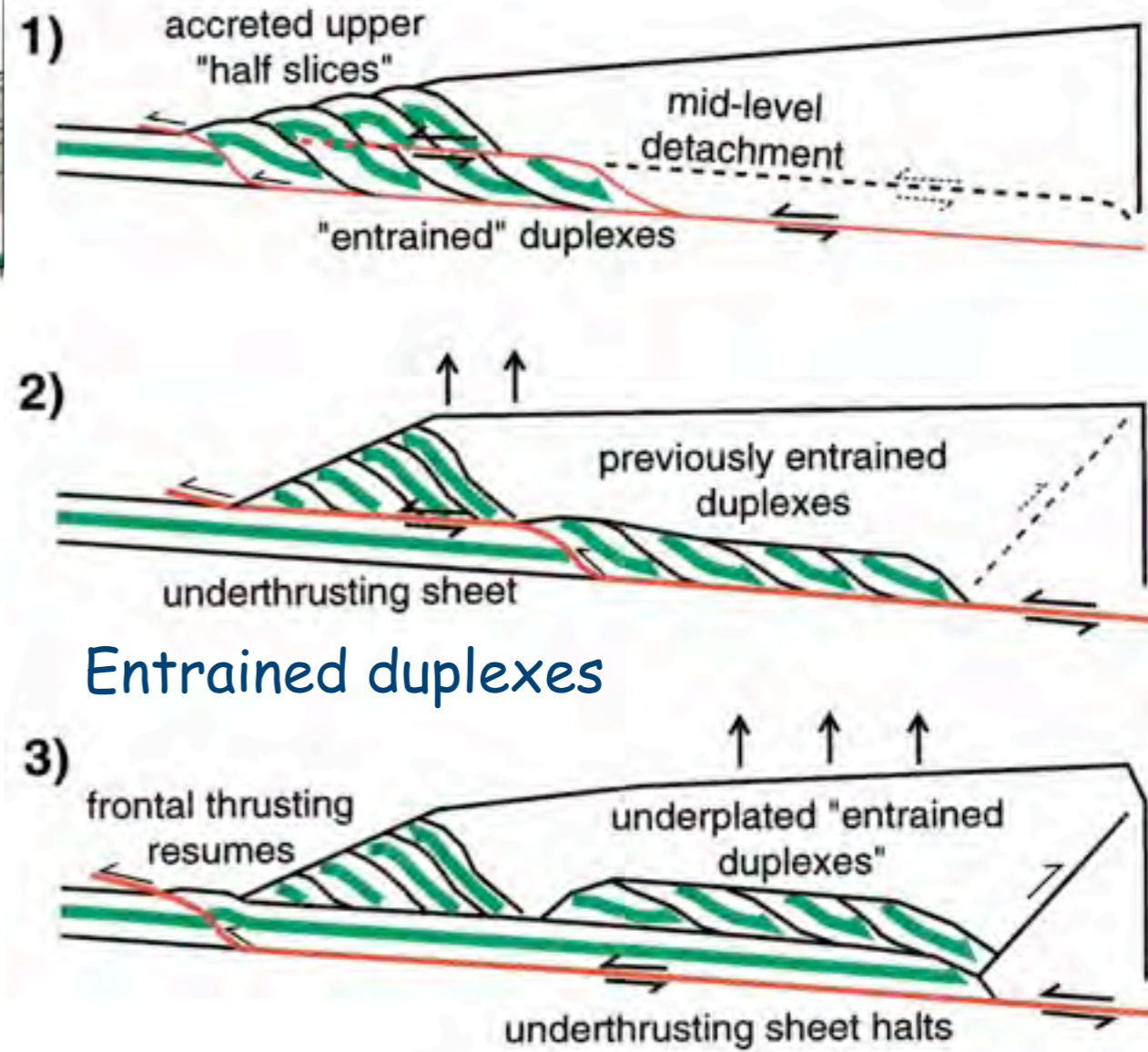
Changes in detachment levels

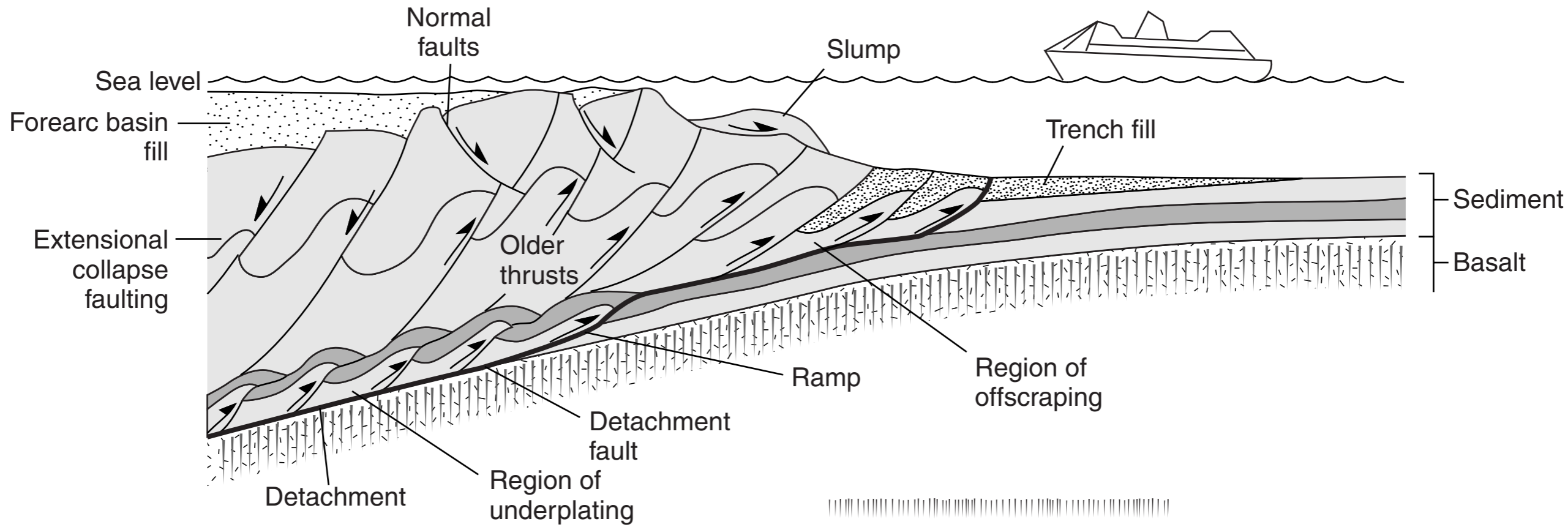


Kukowski, Lallemand, Malavieille, Gutscher & Reston, 2002, Marine Geology

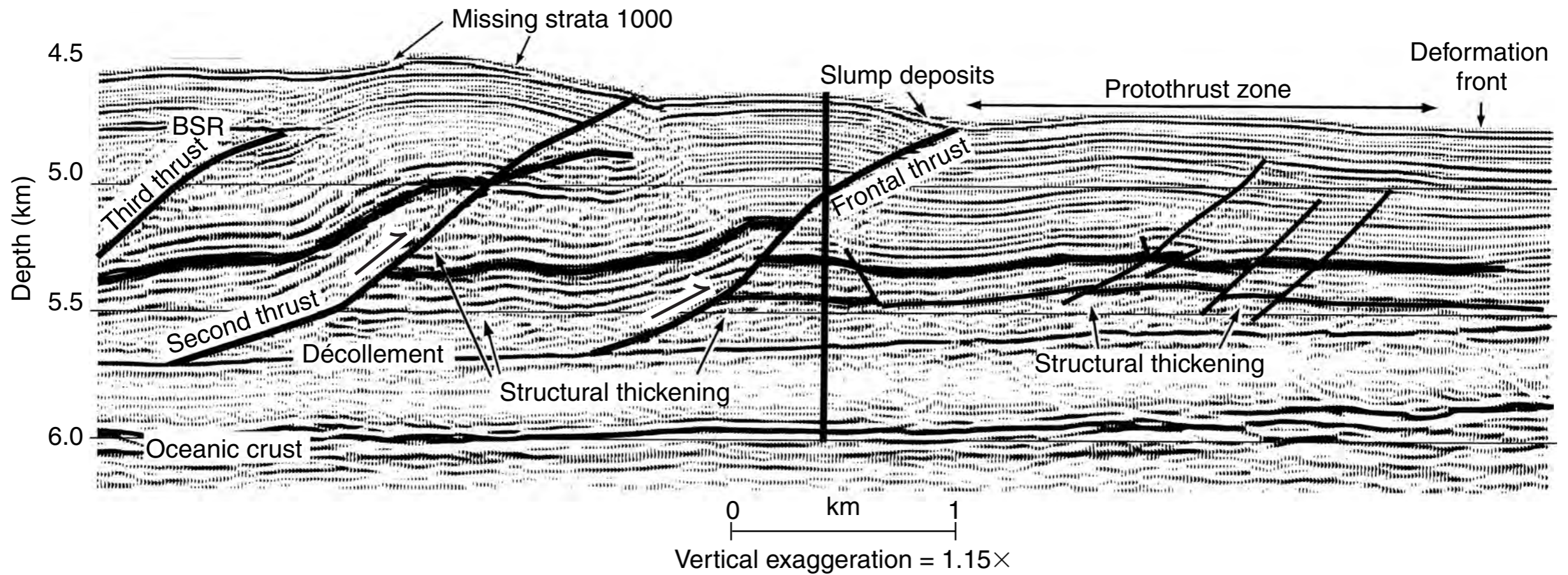
Underplating + Frontal accretion

- Strain partitioning
- Two different growth processes acting simultaneously



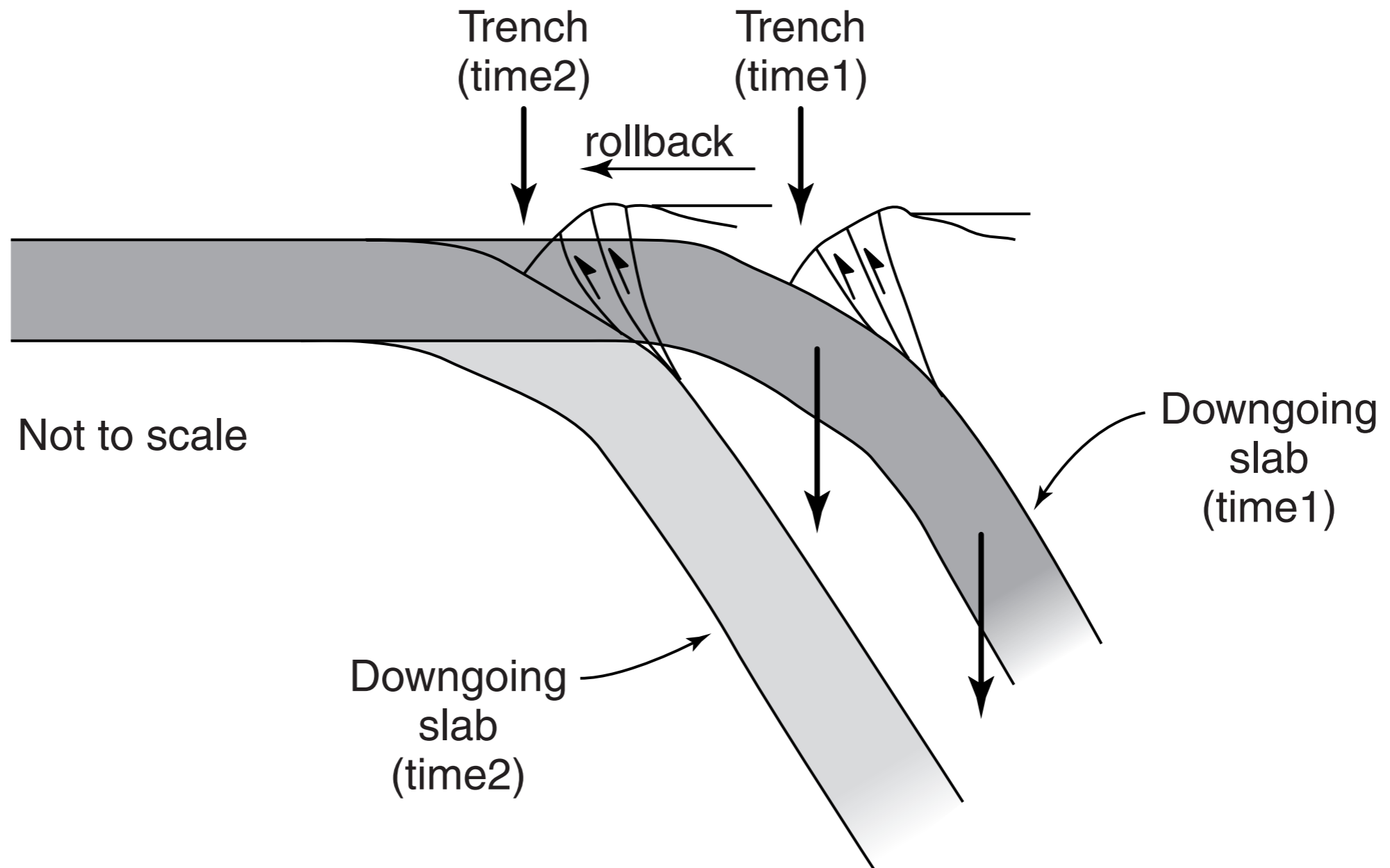


Schematic detail of an accretionary prism, showing different regimes of deformation.



Interpreted seismic-reflection profile of the toe edge of an accretionary prism forming in the Nankai trough off Japan. Several faults can be imaged.

The concept of rollback

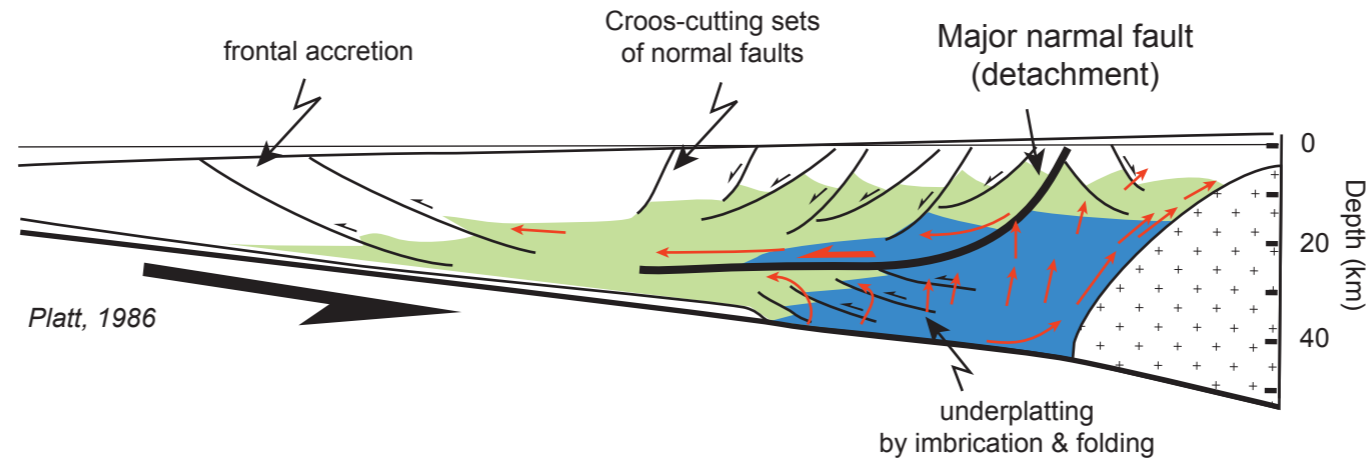


As a subducting slab sinks into the asthenosphere, the position of the trench relative to a fixed point in the mantle migrates. This movement is called rollback

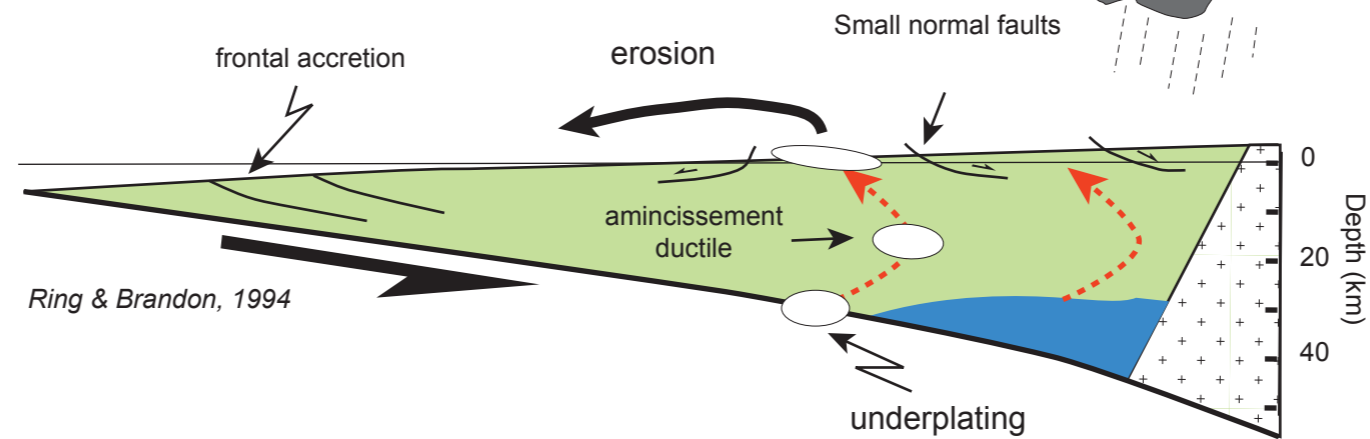
Types of wedge & exhumation of HP rocks

Blueschist

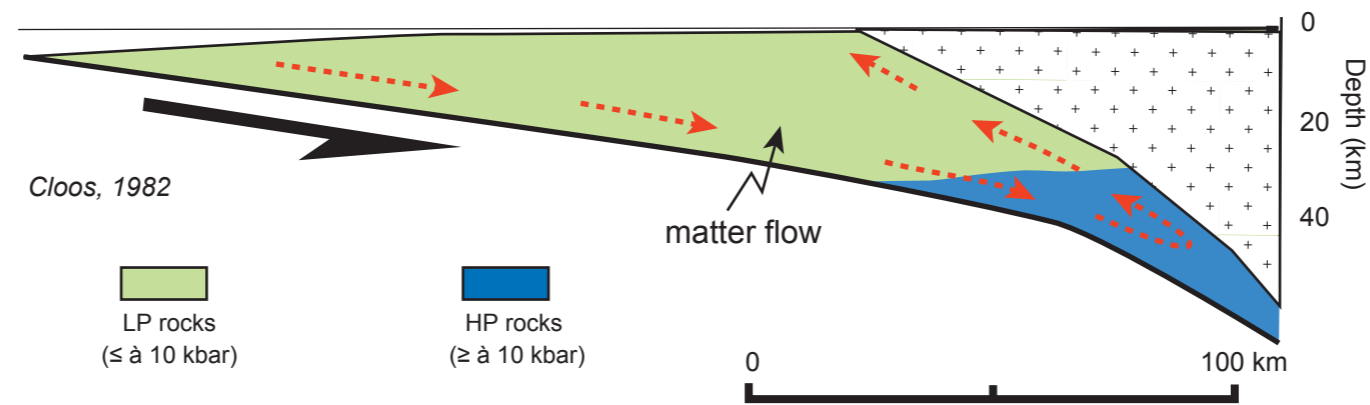
Extension: Crete



Underplating & erosion : South-Chile, Taiwan



"Corner flow" or subduction channel : Western Alps



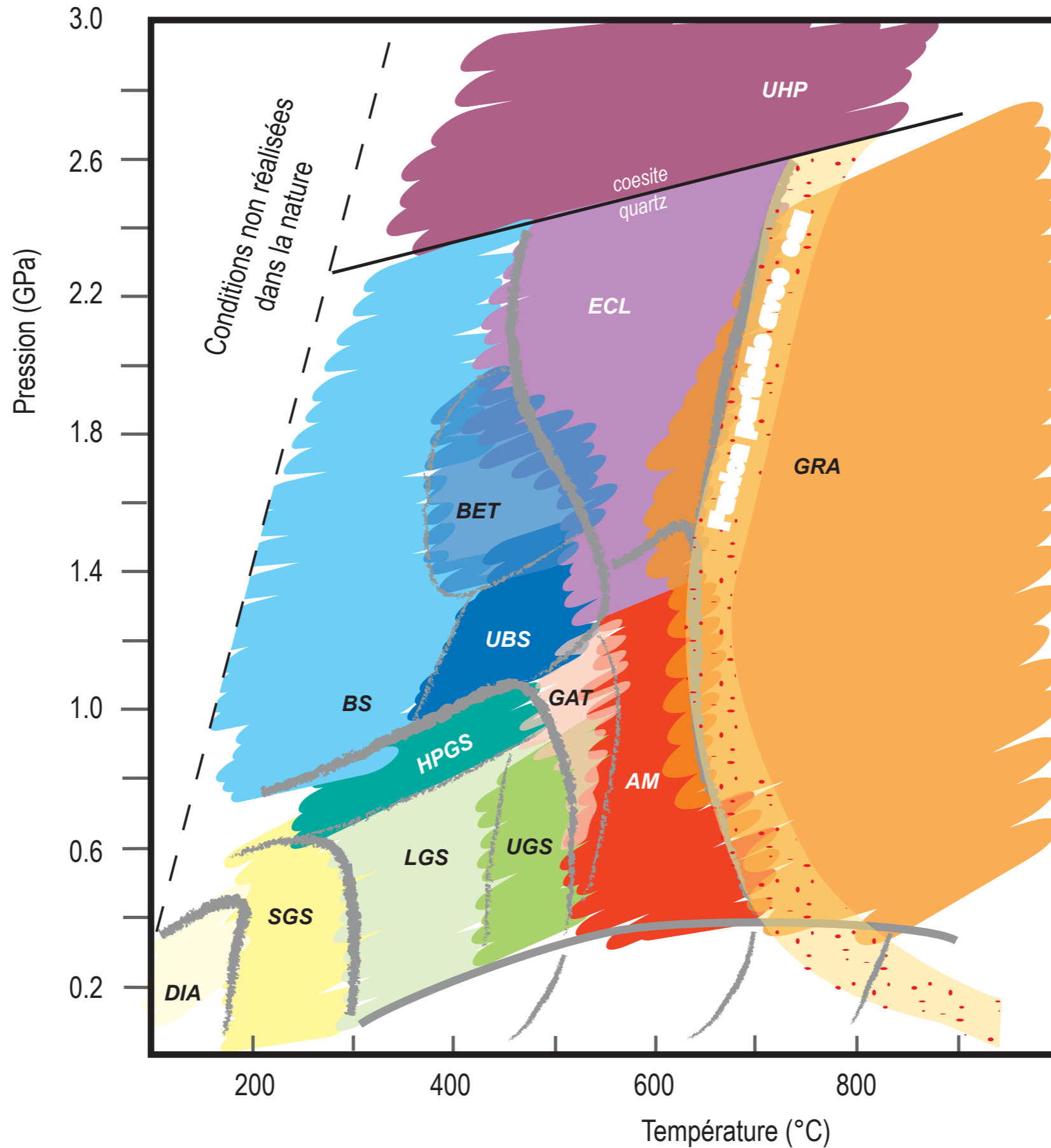
LP rocks
(\leq à 10 kbar)

HP rocks
(\geq à 10 kbar)

after Bousquet, 1998

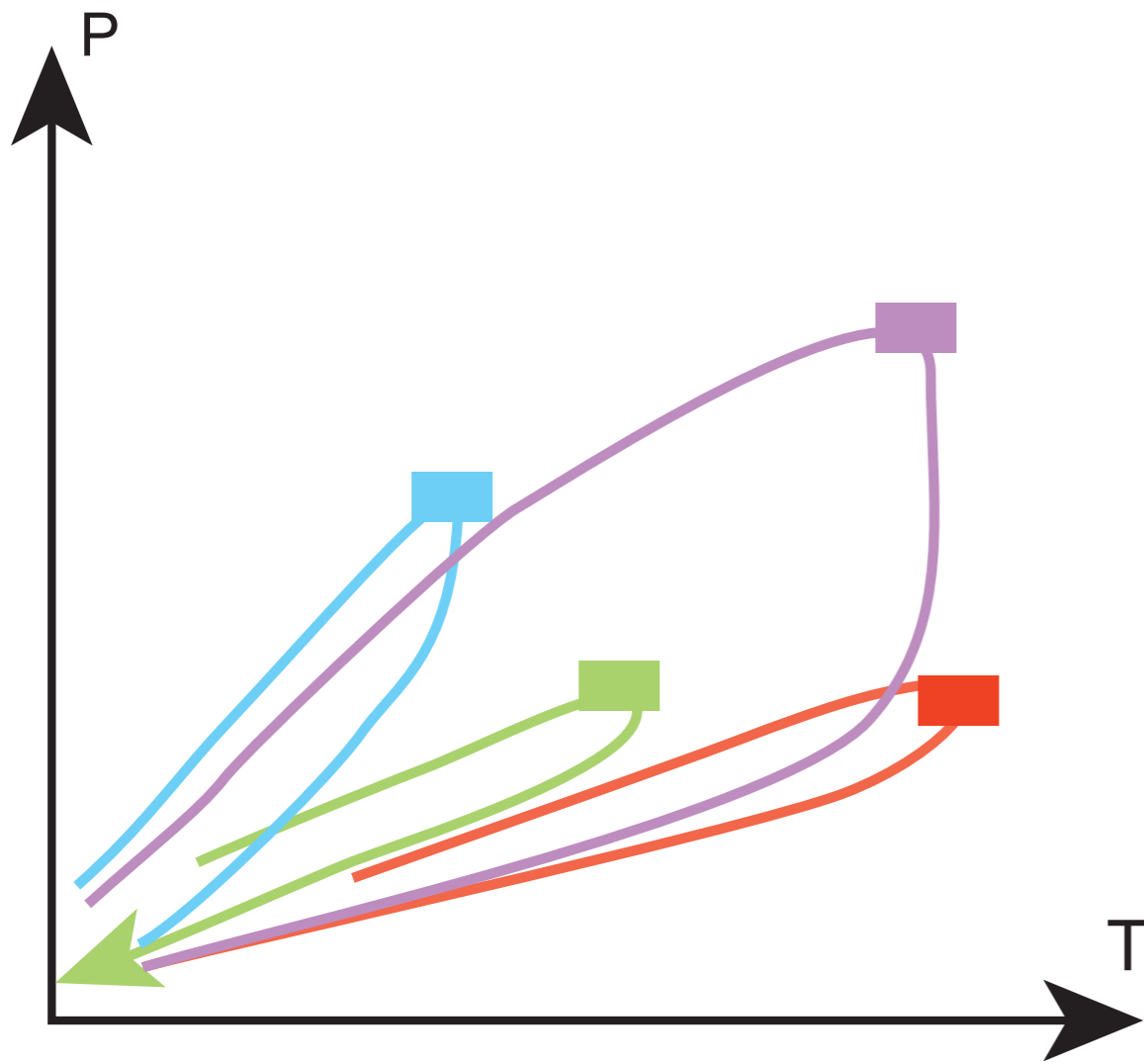
Blueschist, Eclogites

Metamorphic facies

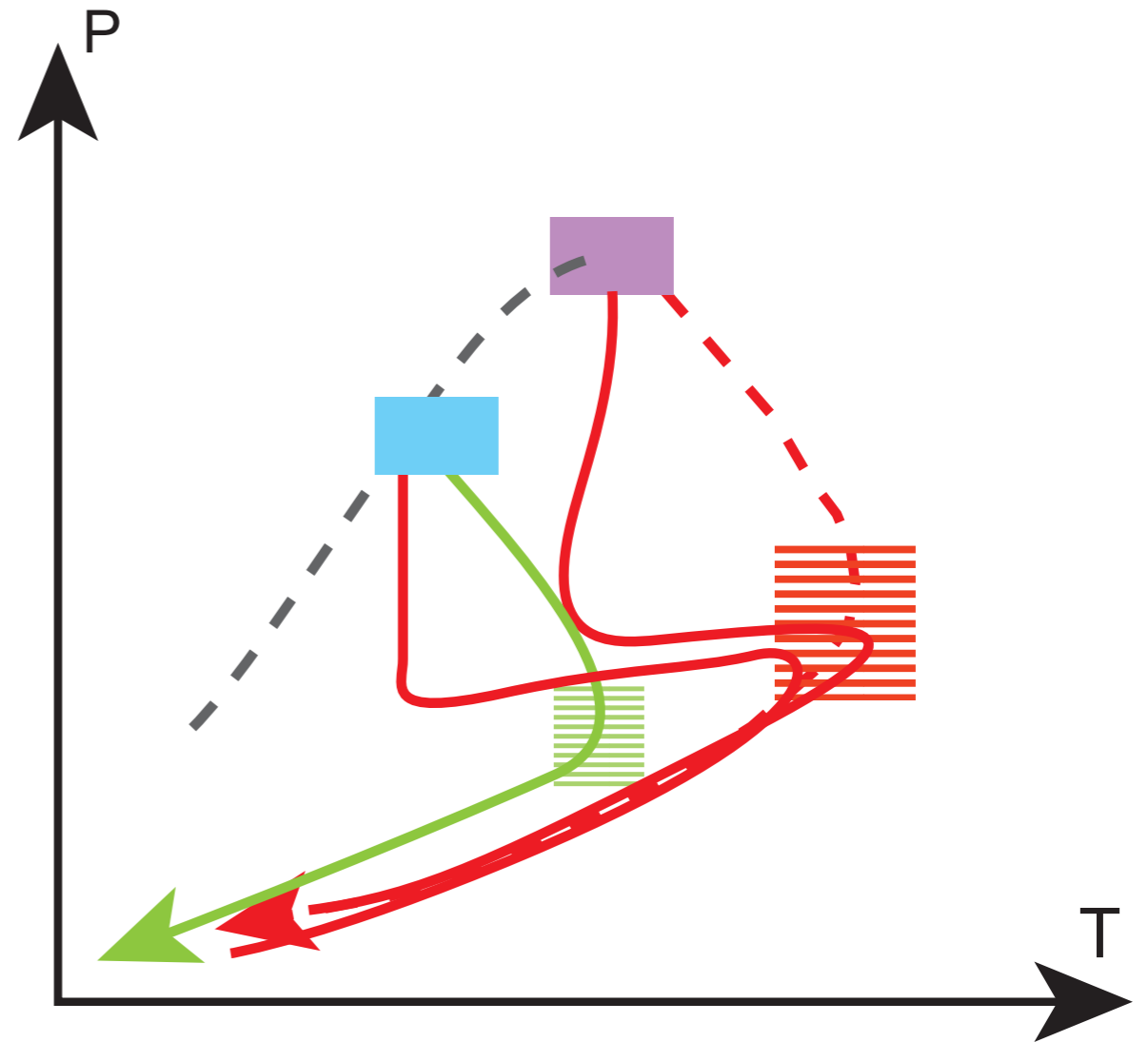


Different types of PT paths

Single metamorphic peak always followed by cooling or isothermal retrograde events



Metamorphic P-peak overprinted by a later thermal peak

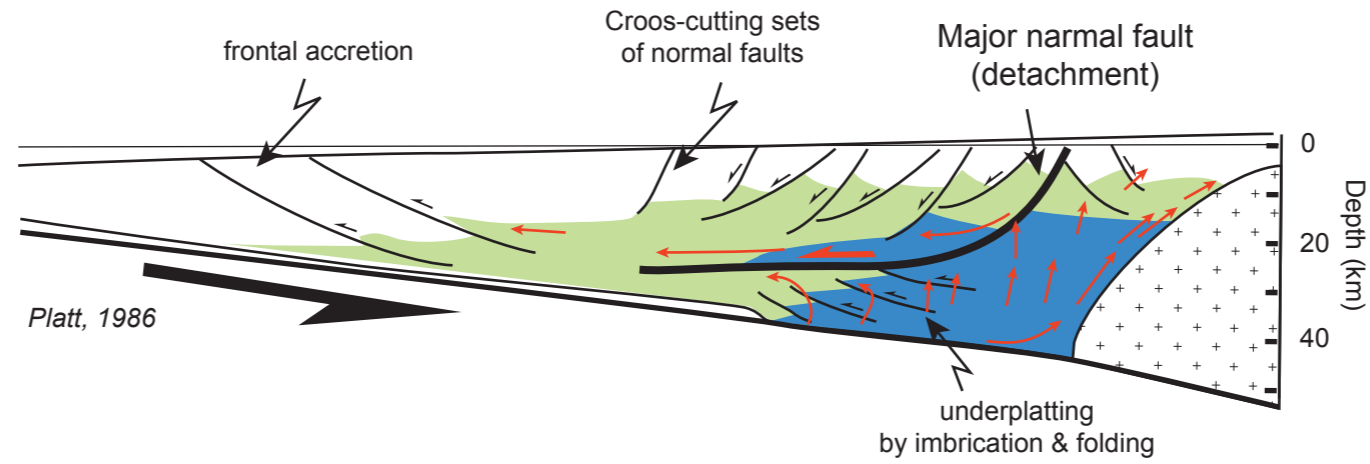


Oberhänsli et al., 2004

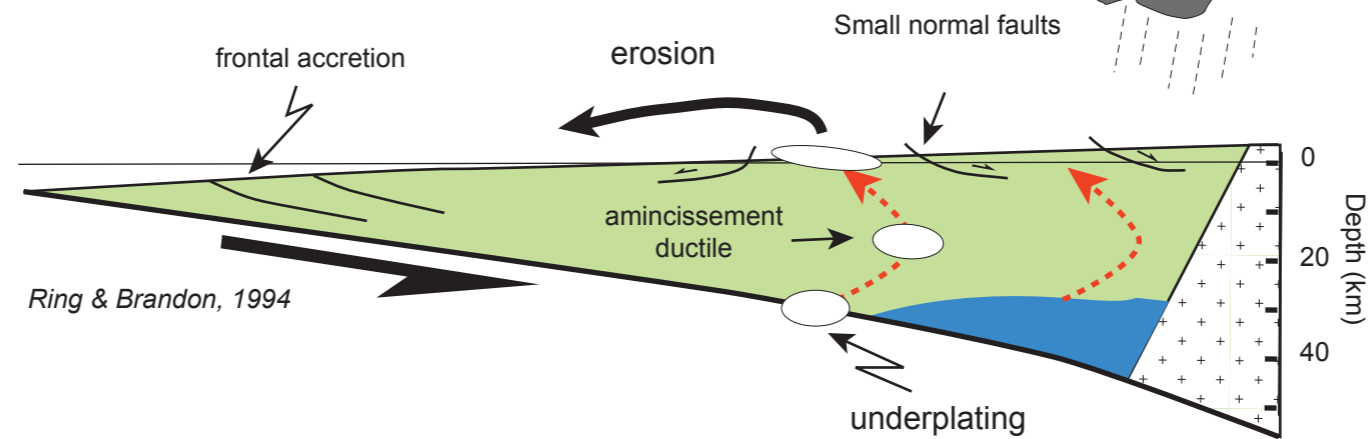
Types of wedge & exhumation of HP rocks

Blueschist

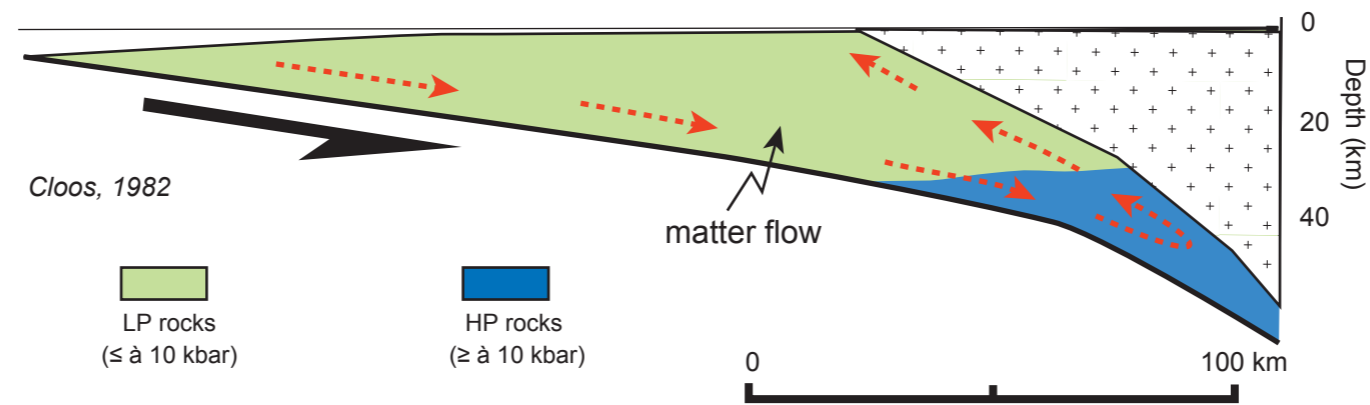
Extension: Crete



Underplating & erosion : South-Chile, Taiwan



"Corner flow" or subduction channel : Western Alps



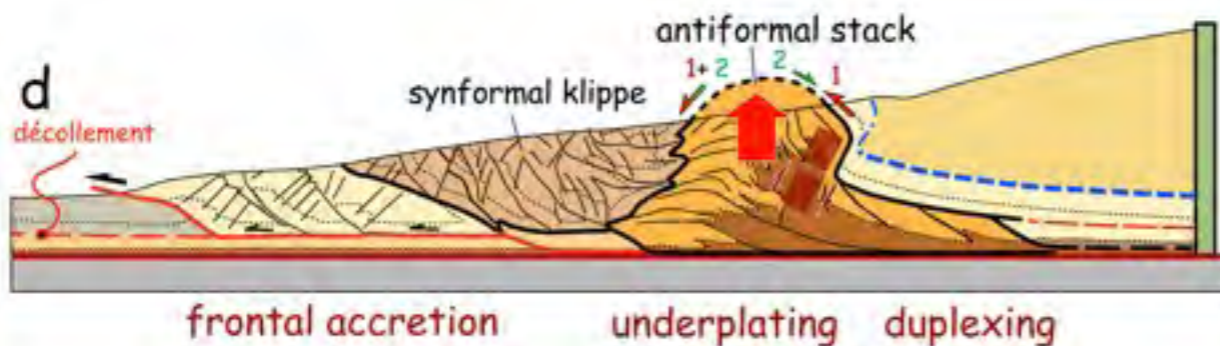
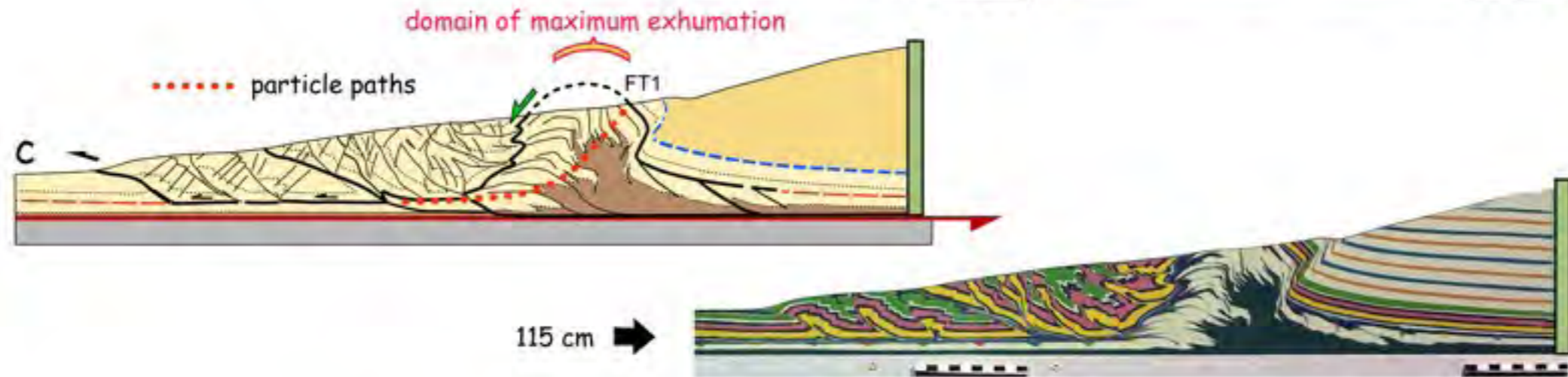
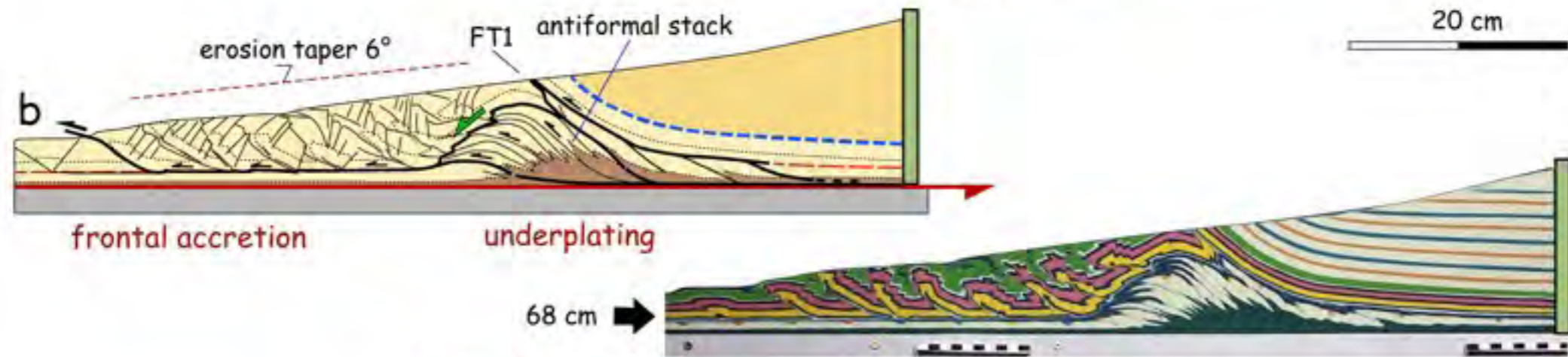
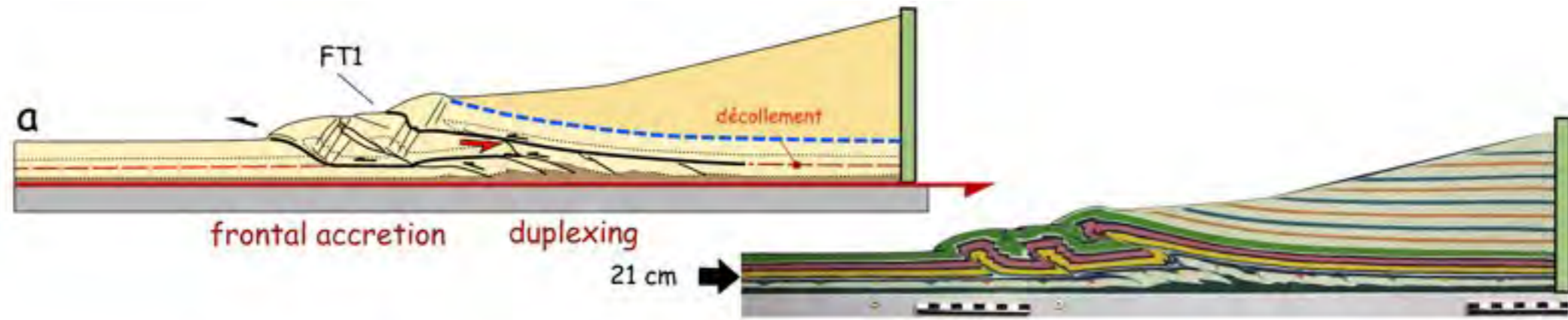
LP rocks
(\leq à 10 kbar)

HP rocks
(\geq à 10 kbar)

after Bousquet, 1998

Blueschist, Eclogites

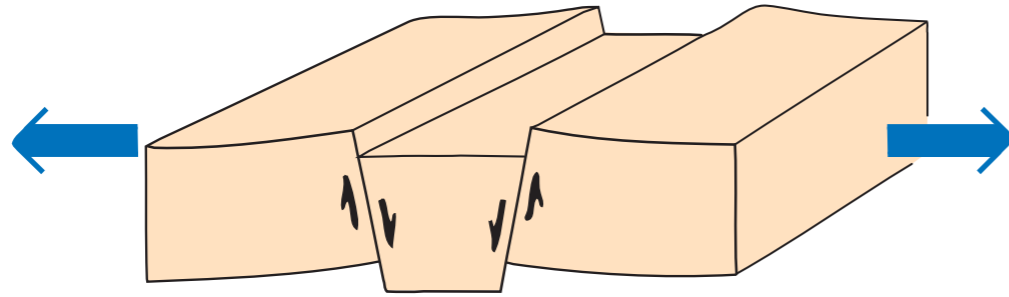
Accretion & erosion for the exhumation



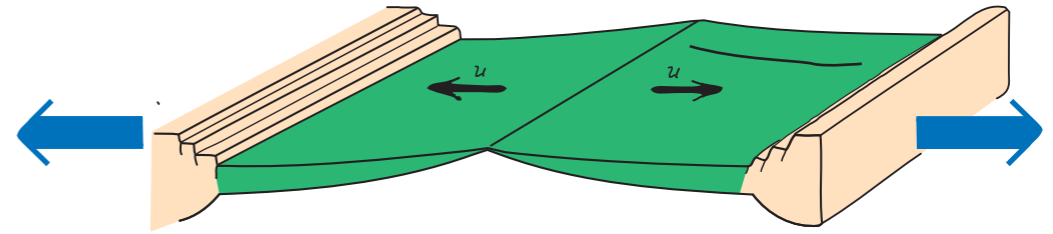
Underplating + erosion
Rapid exhumation

Wilson cycle

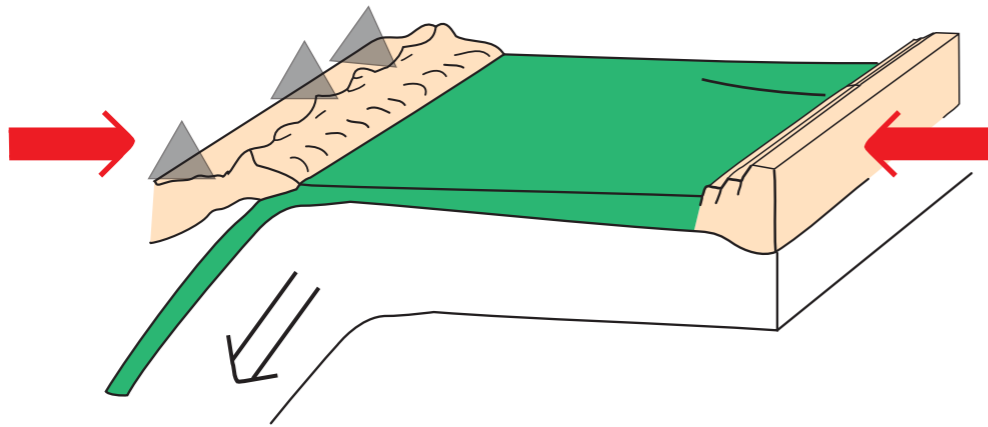
Rifting



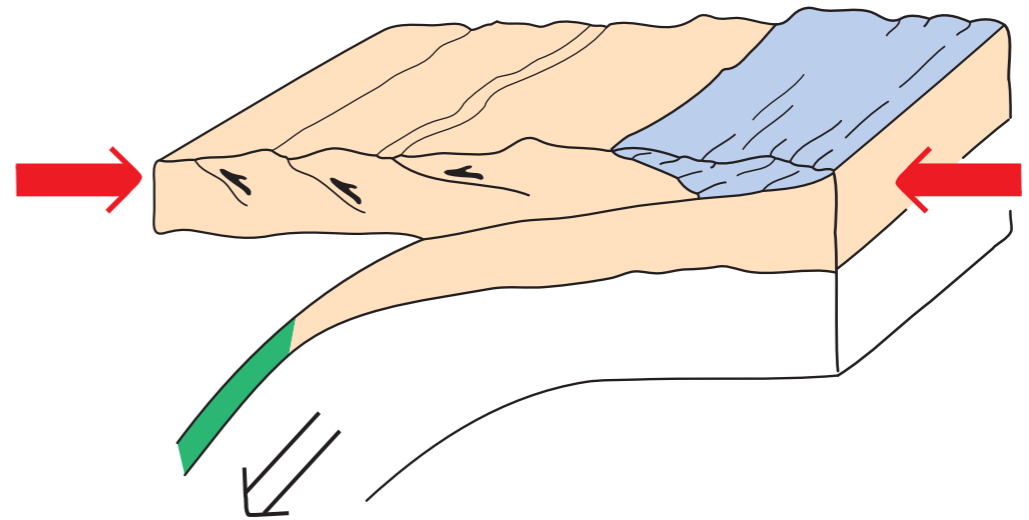
Ocean formation



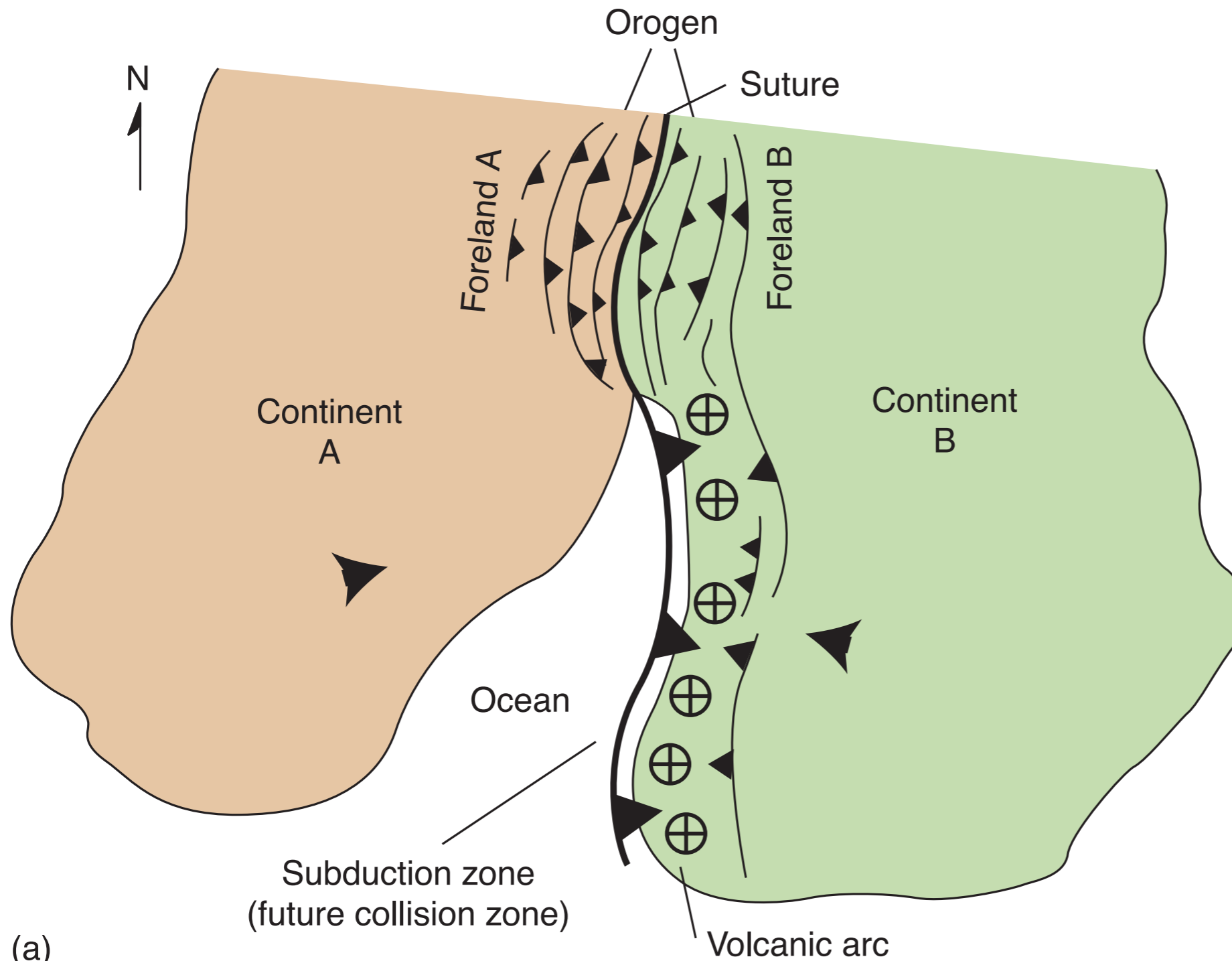
Subduction



Collision



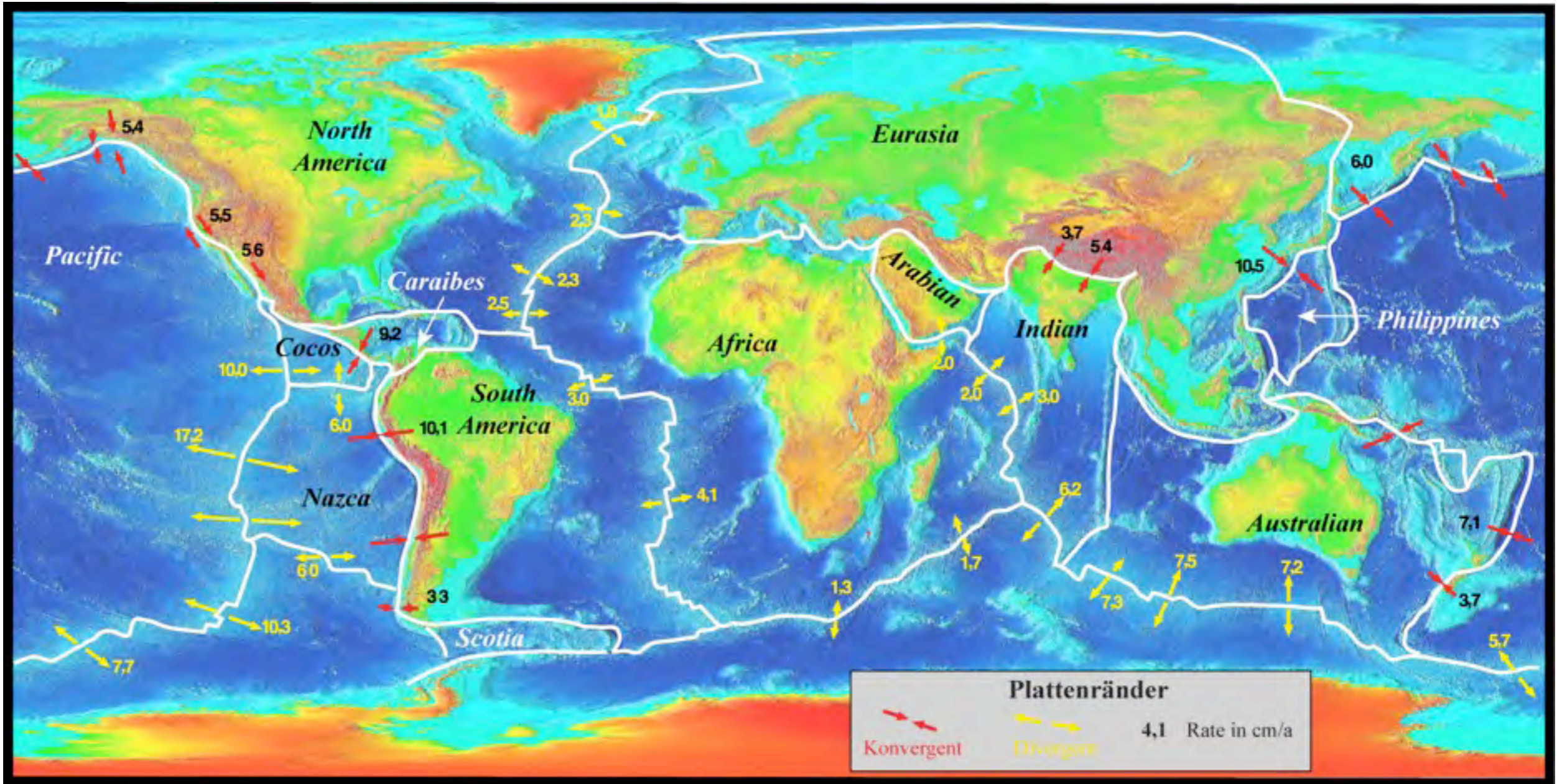
From subduction to collision



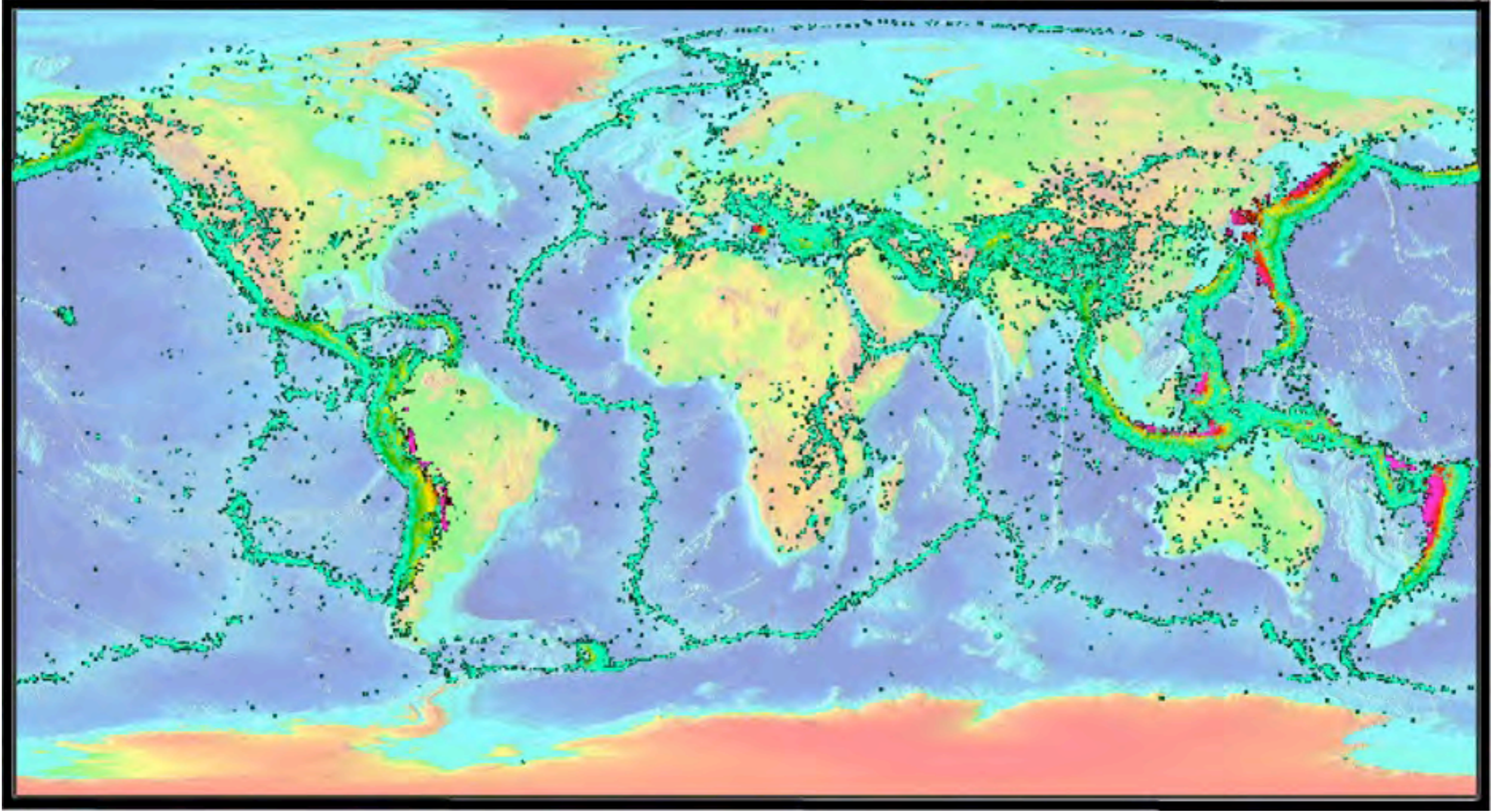
(a)

A map showing a zipper-like collision between two continents. Here, the ocean between the two continents is closing progressively from north to south. In the collision zone, the boundary between what had originally been two separate continents.

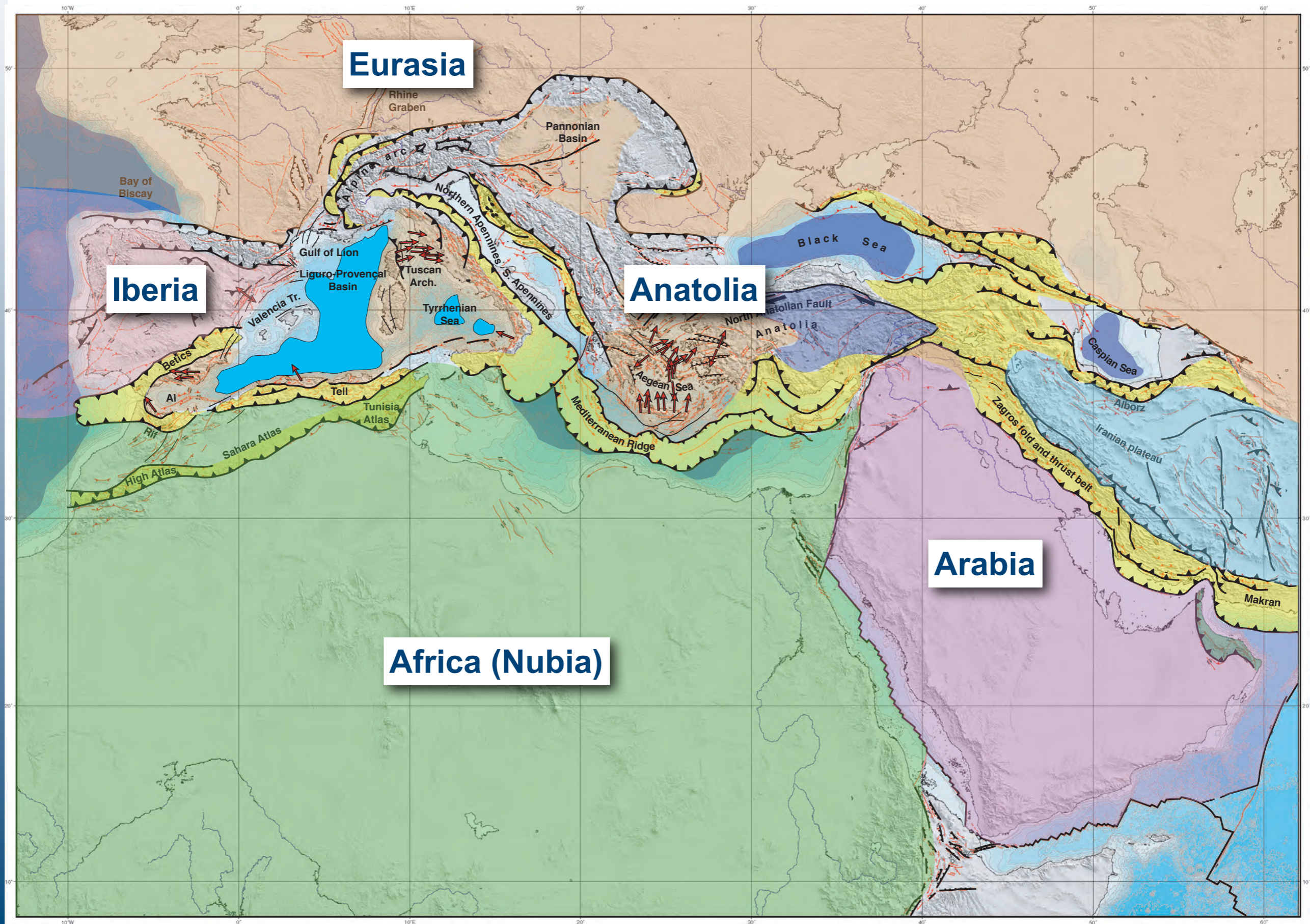
Konvergente Plattenränder



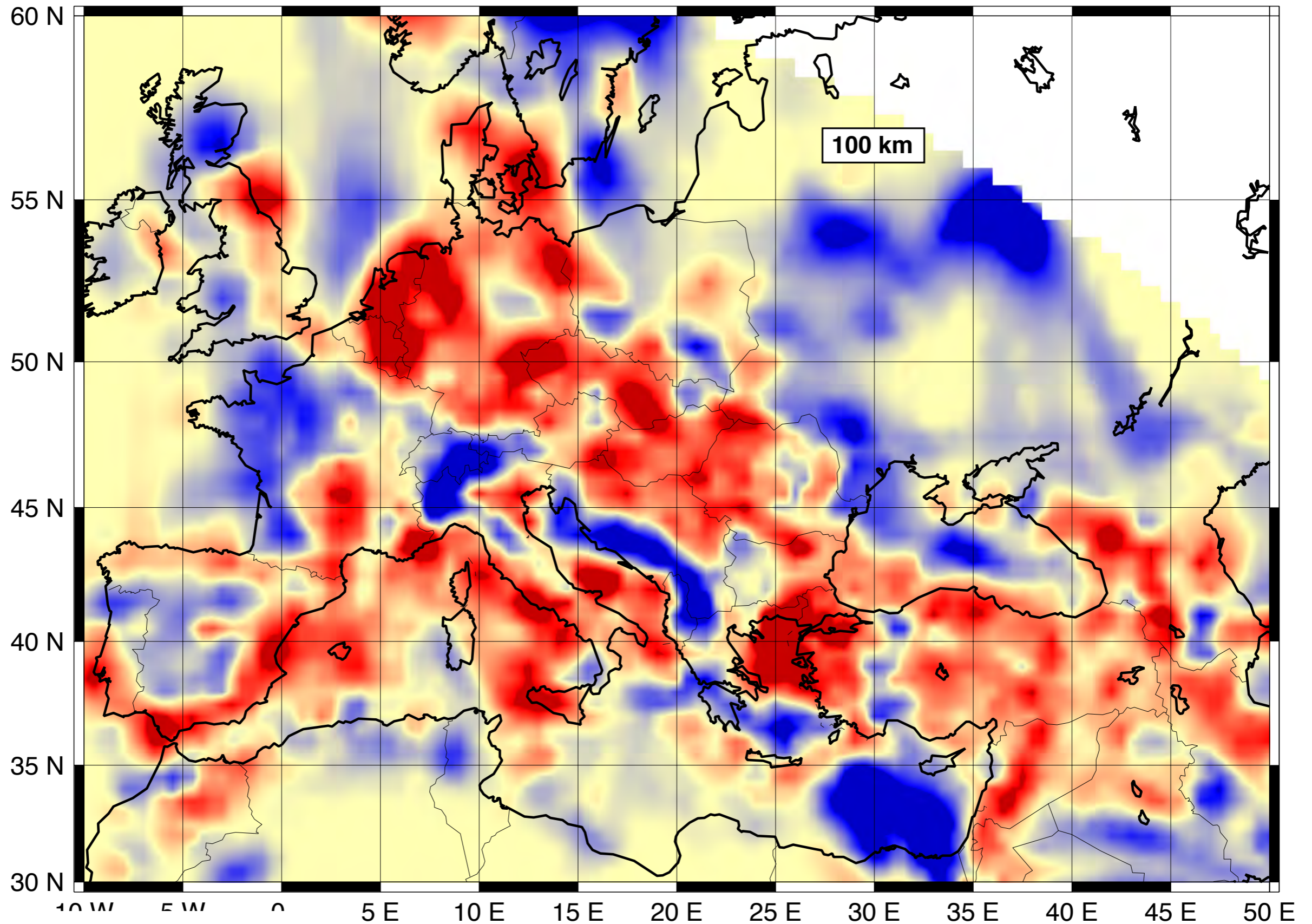
Earthquakes



Deep structure of the Mediterranean area

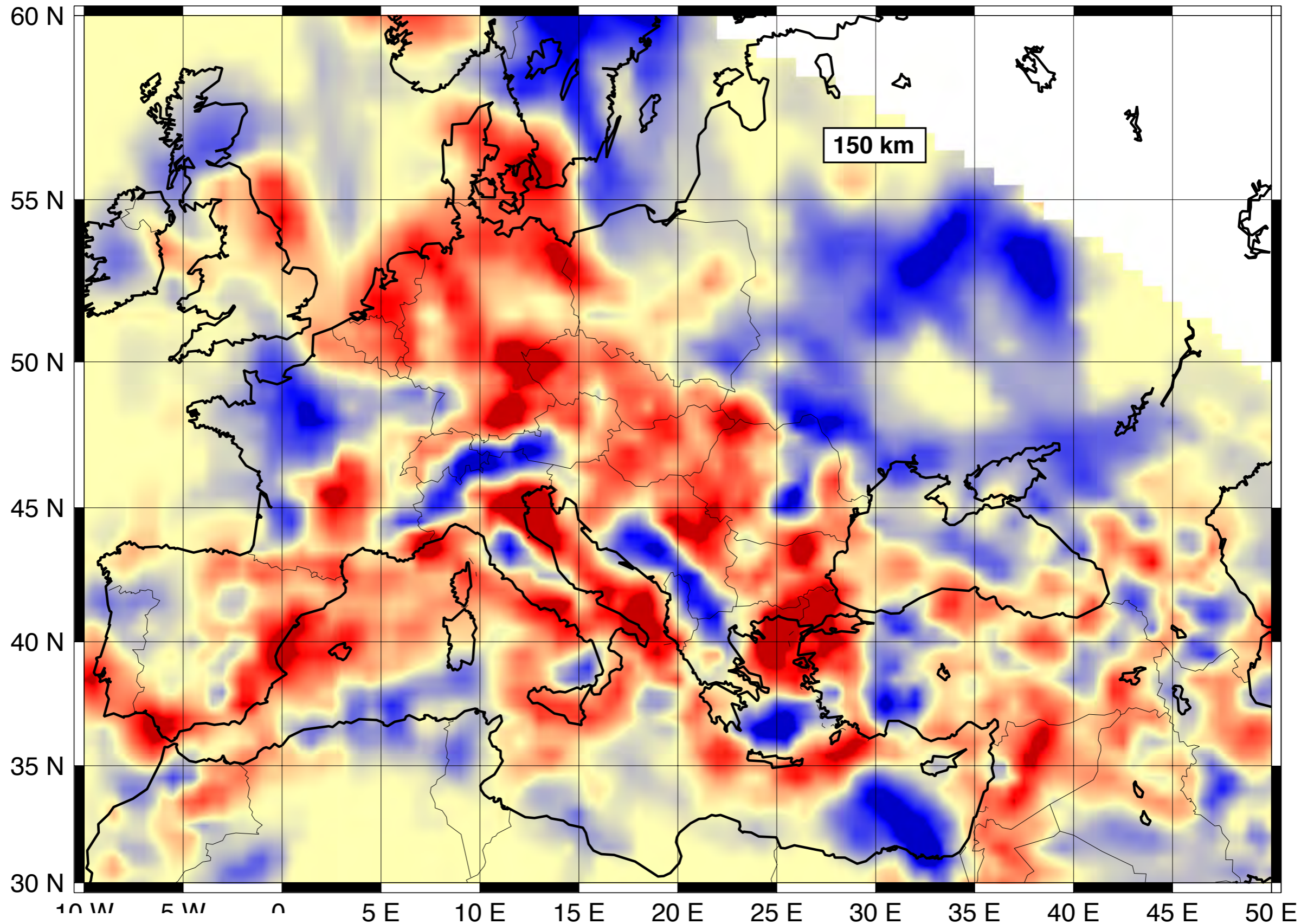


Deep structure of the Mediterranean area

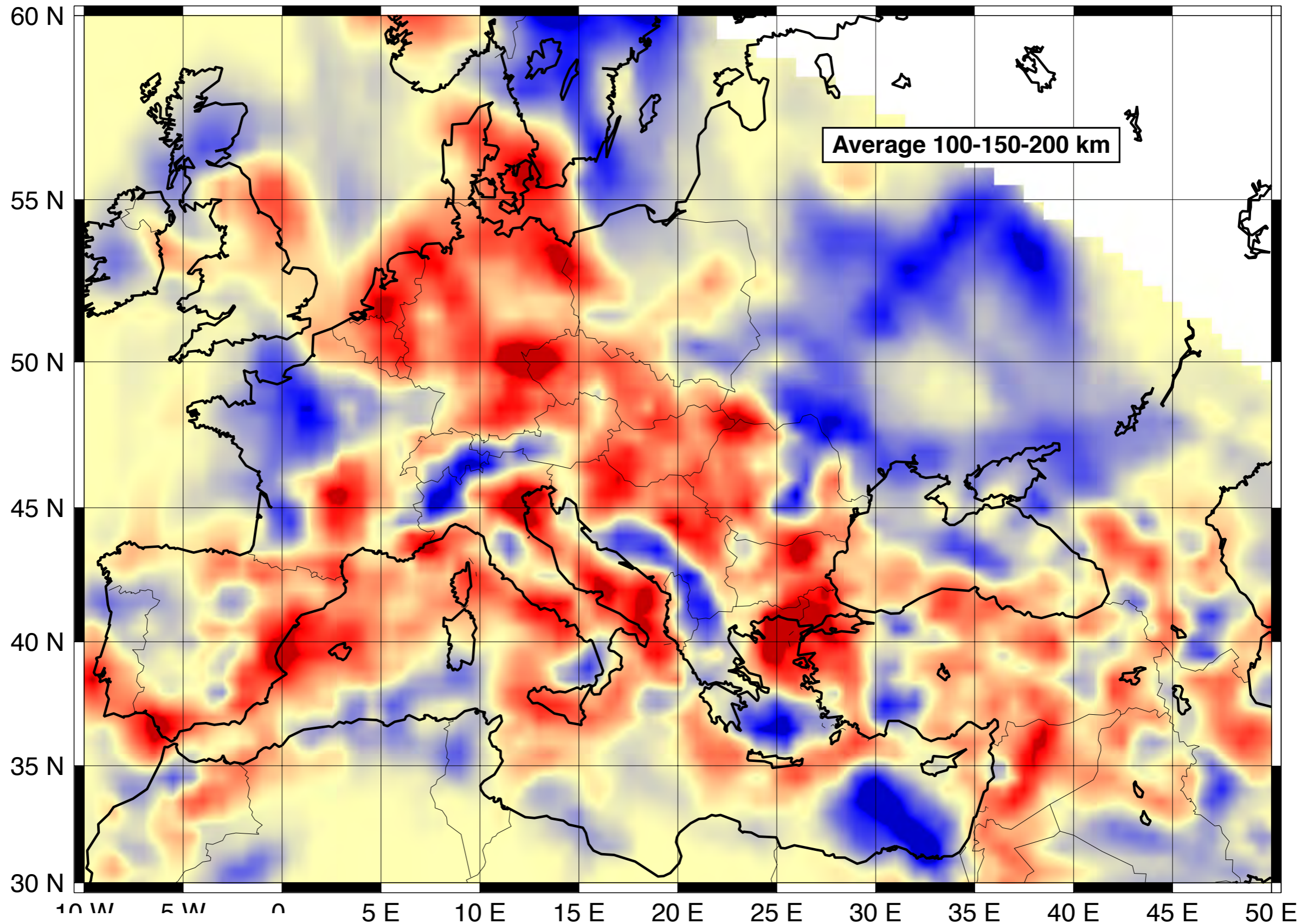


Piromallo et al.

Deep structure of the Mediterranean area

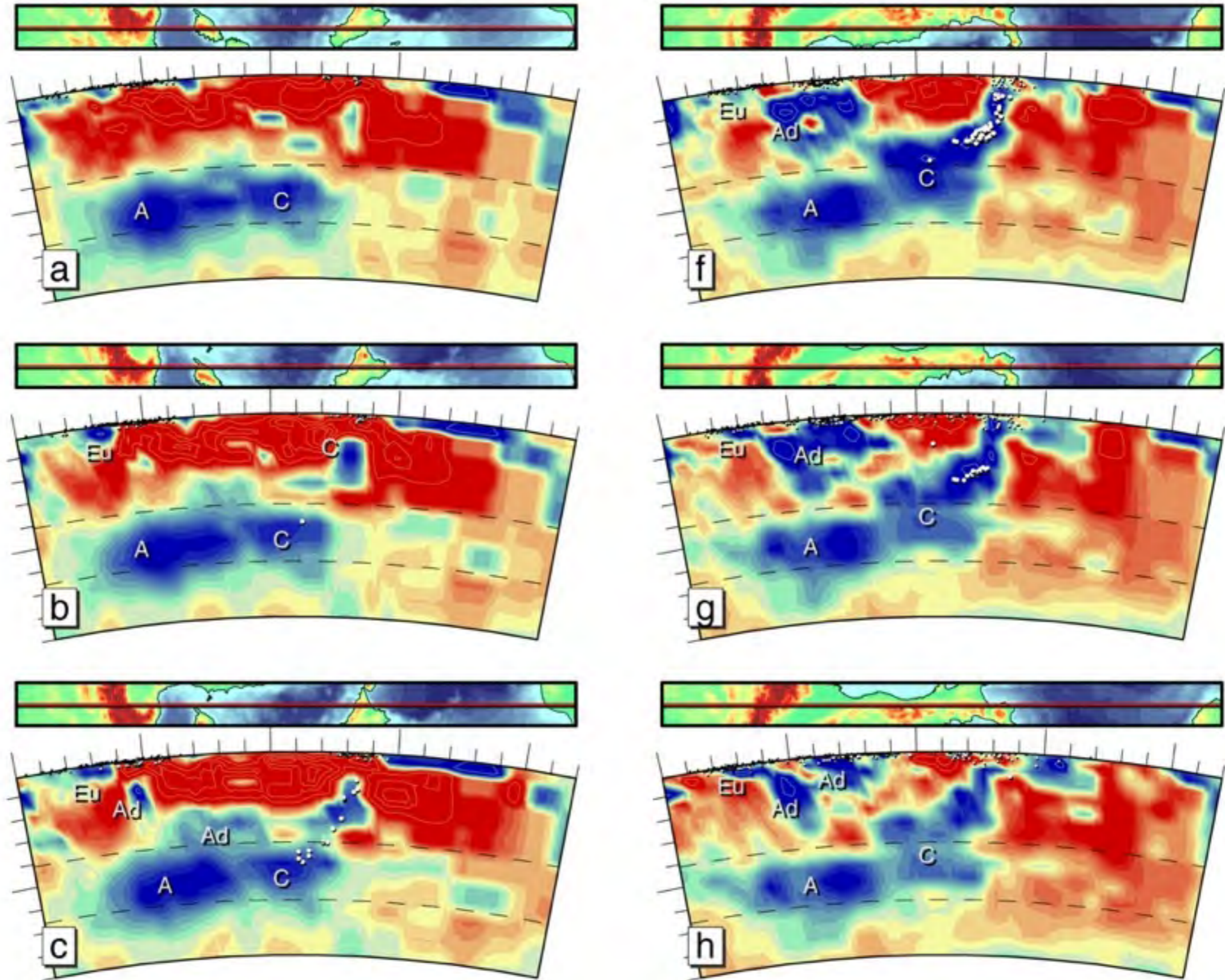


Deep structure of the Mediterranean area

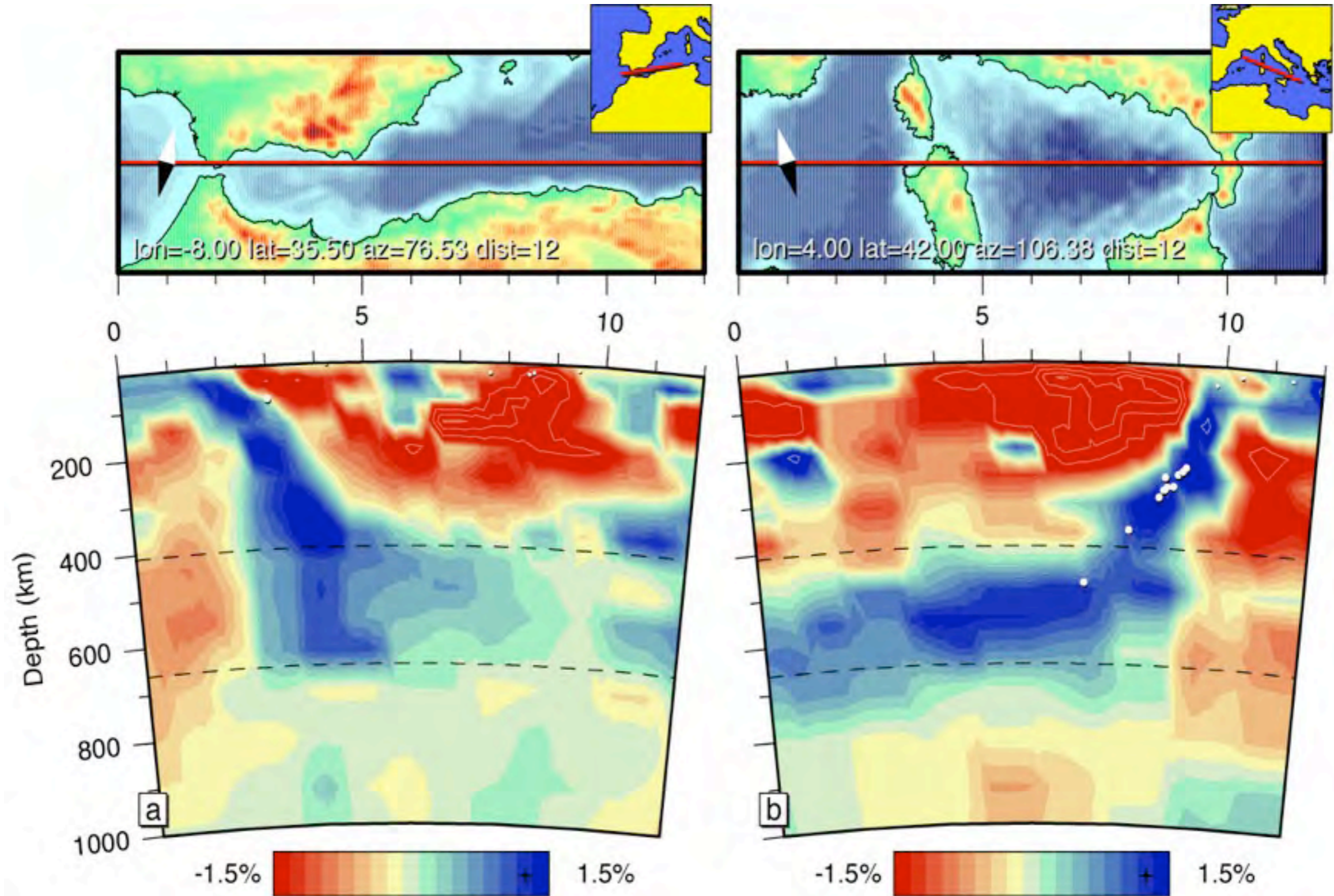


Piromallo et al.

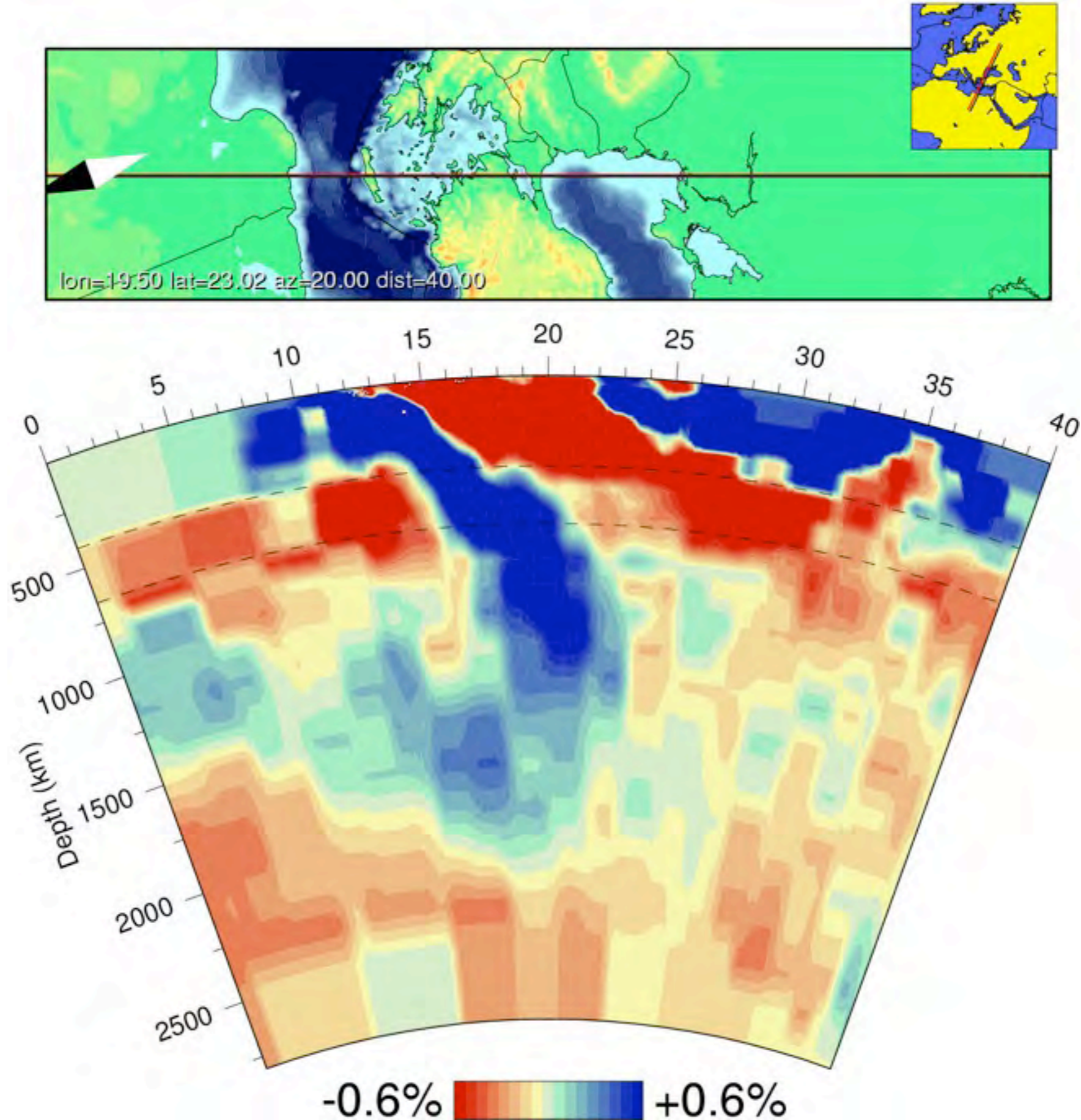
Deep structure of the Mediterranean area



Deep structure of the Mediterranean area

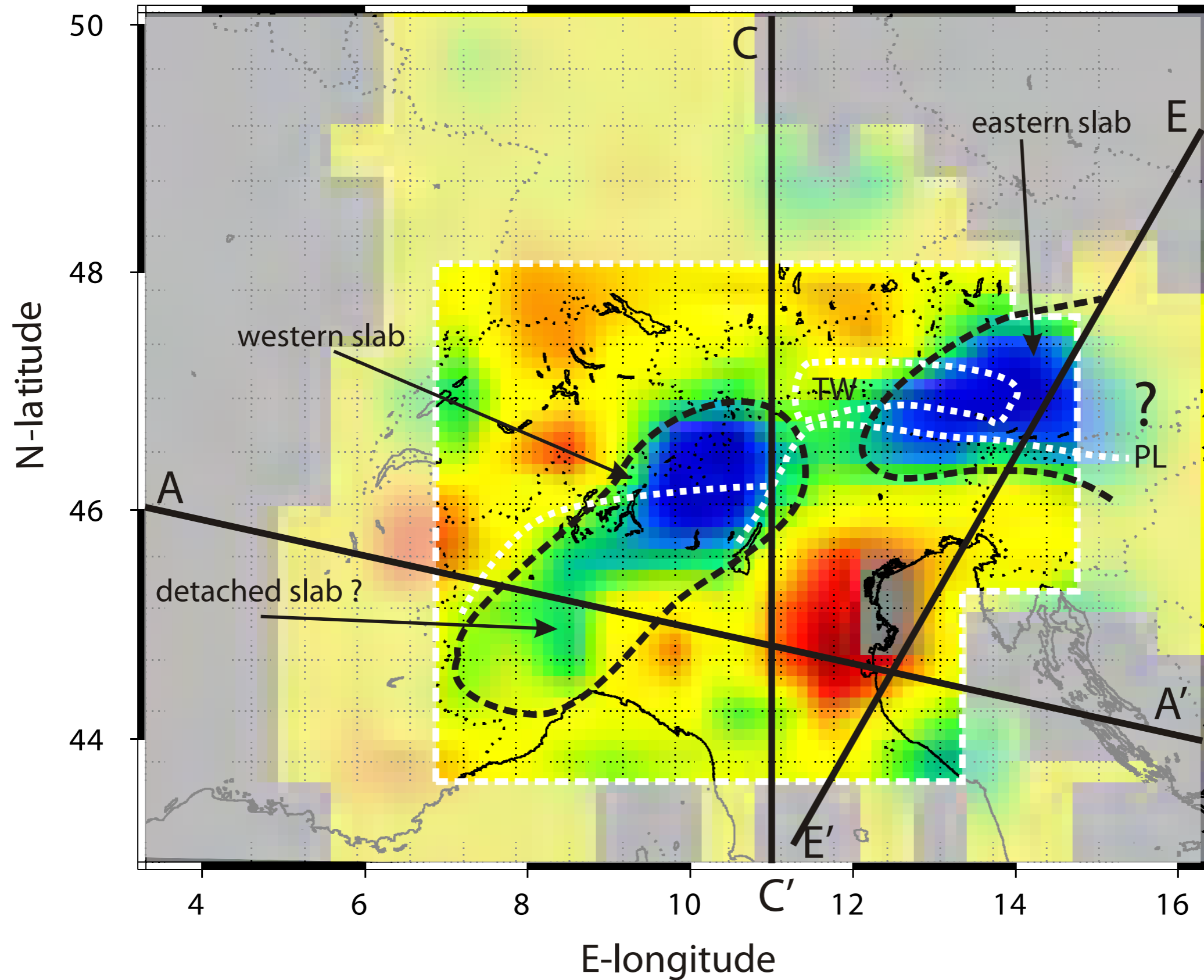


Deep structure of the Mediterranean area

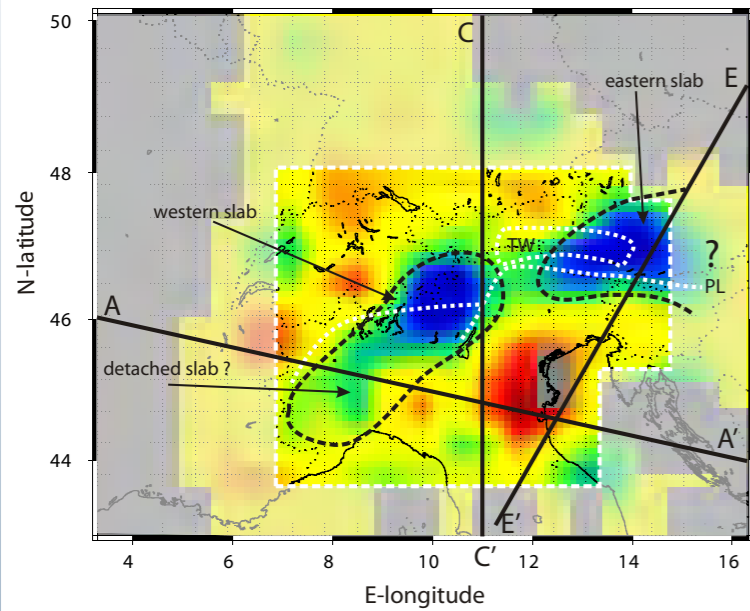


Spakman...

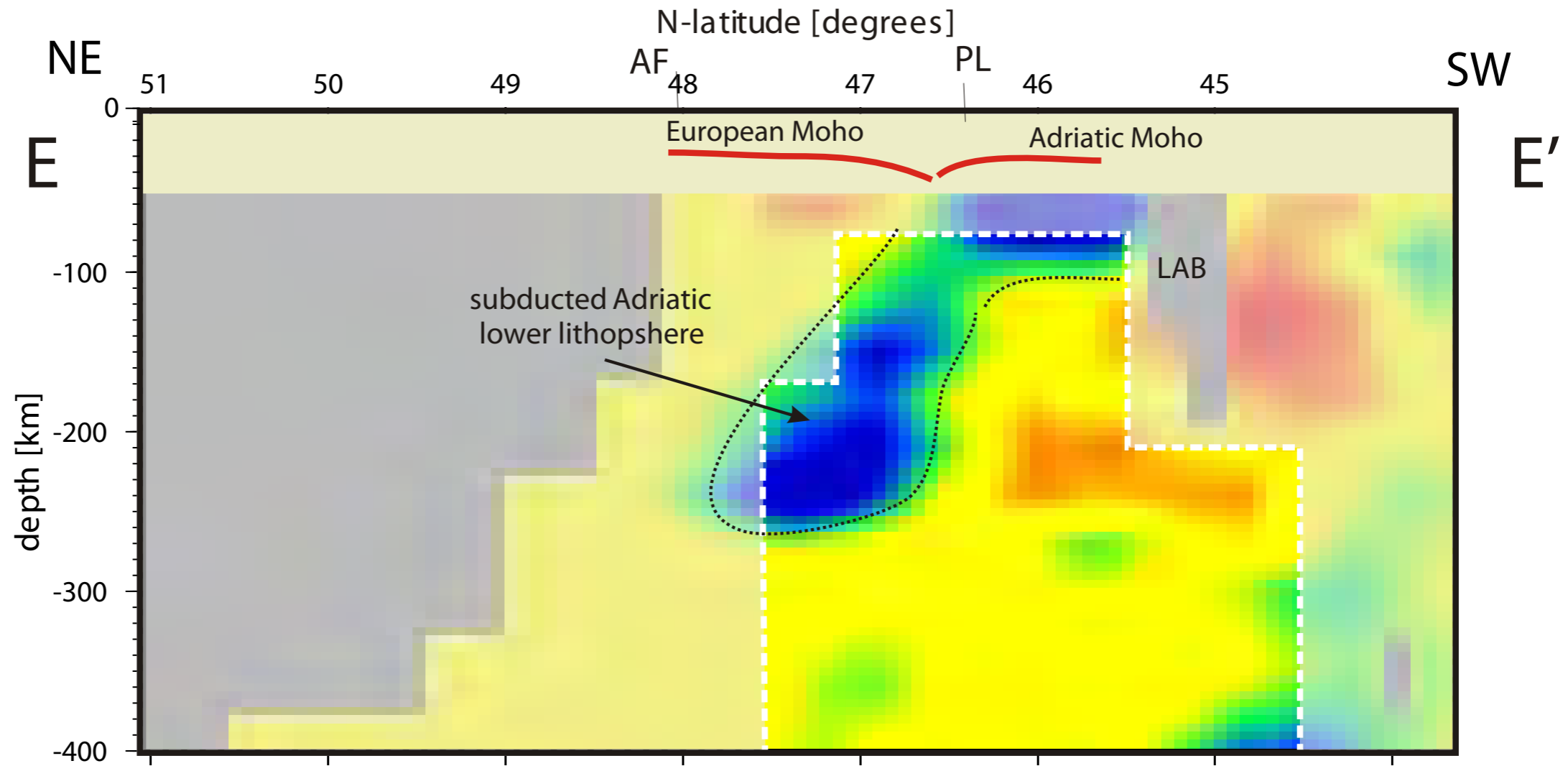
Deep structure of the Alpine area



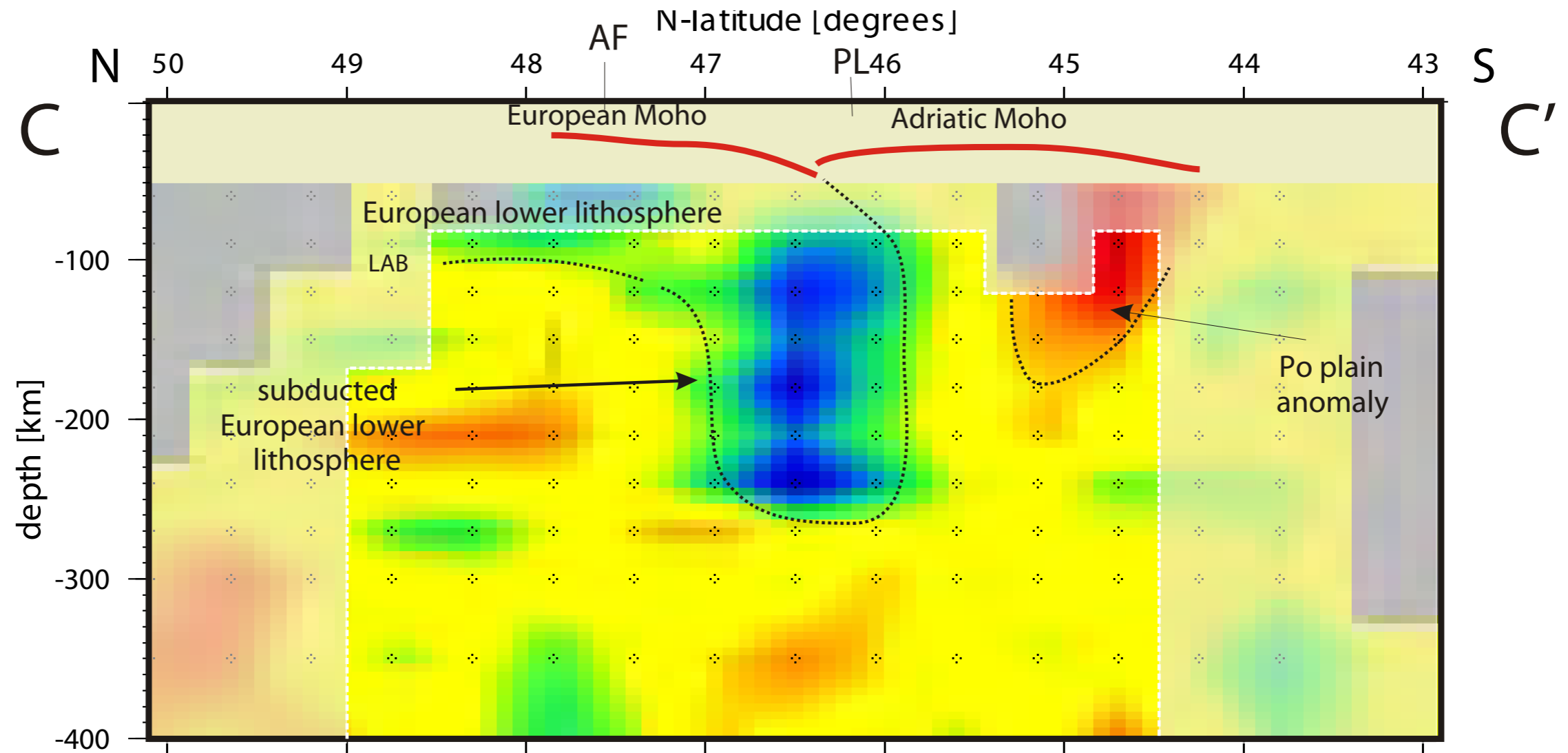
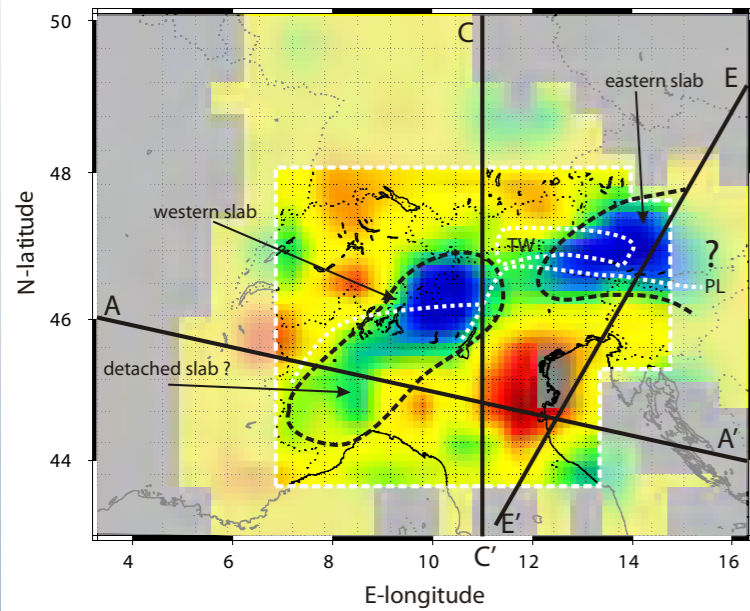
Deep structure of the Alpine area: East



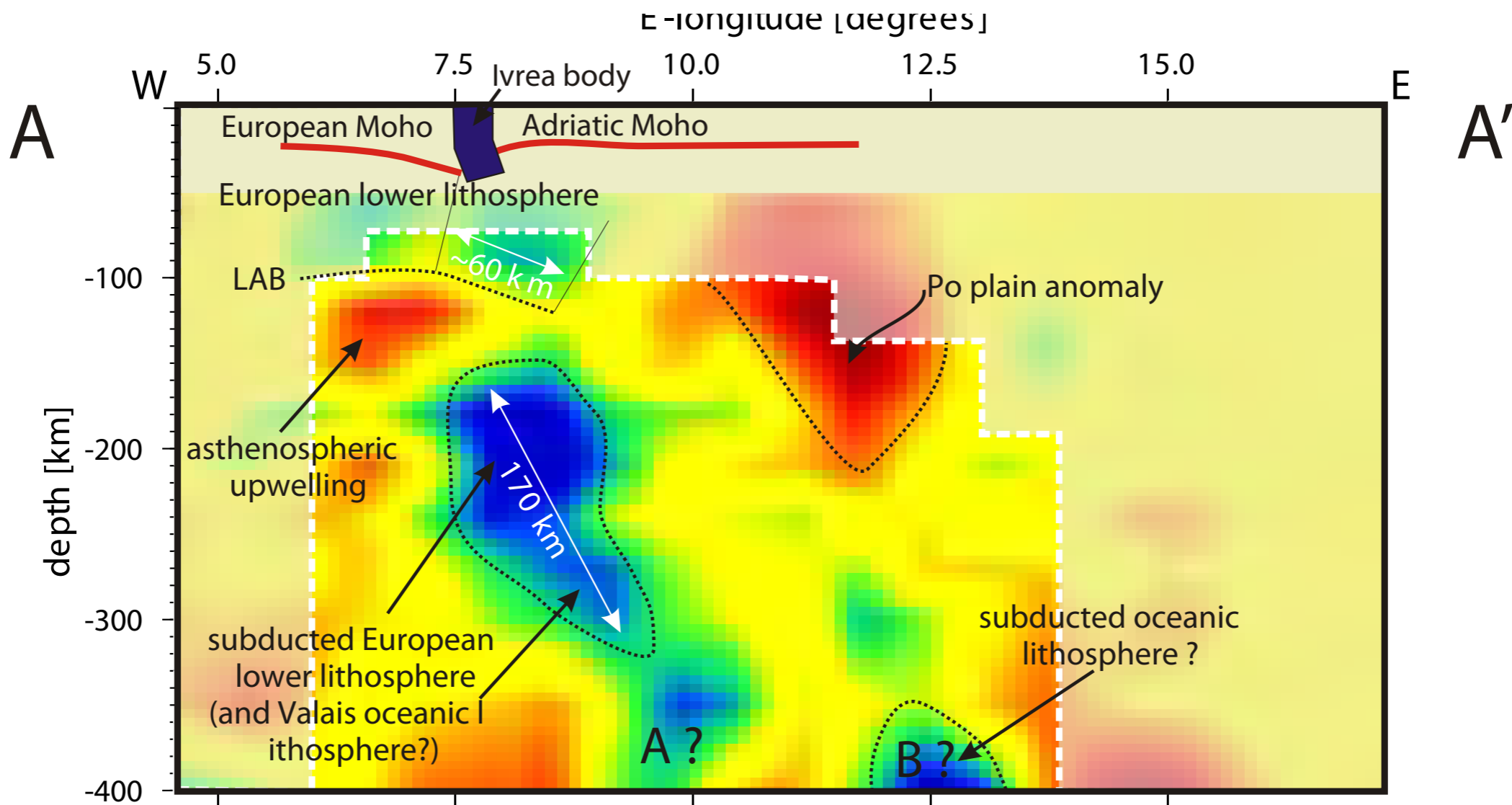
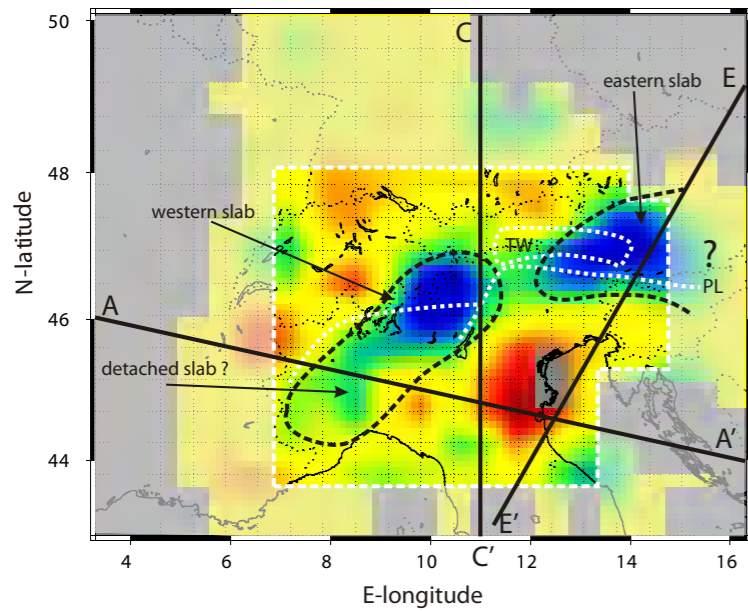
f) Profile EE'



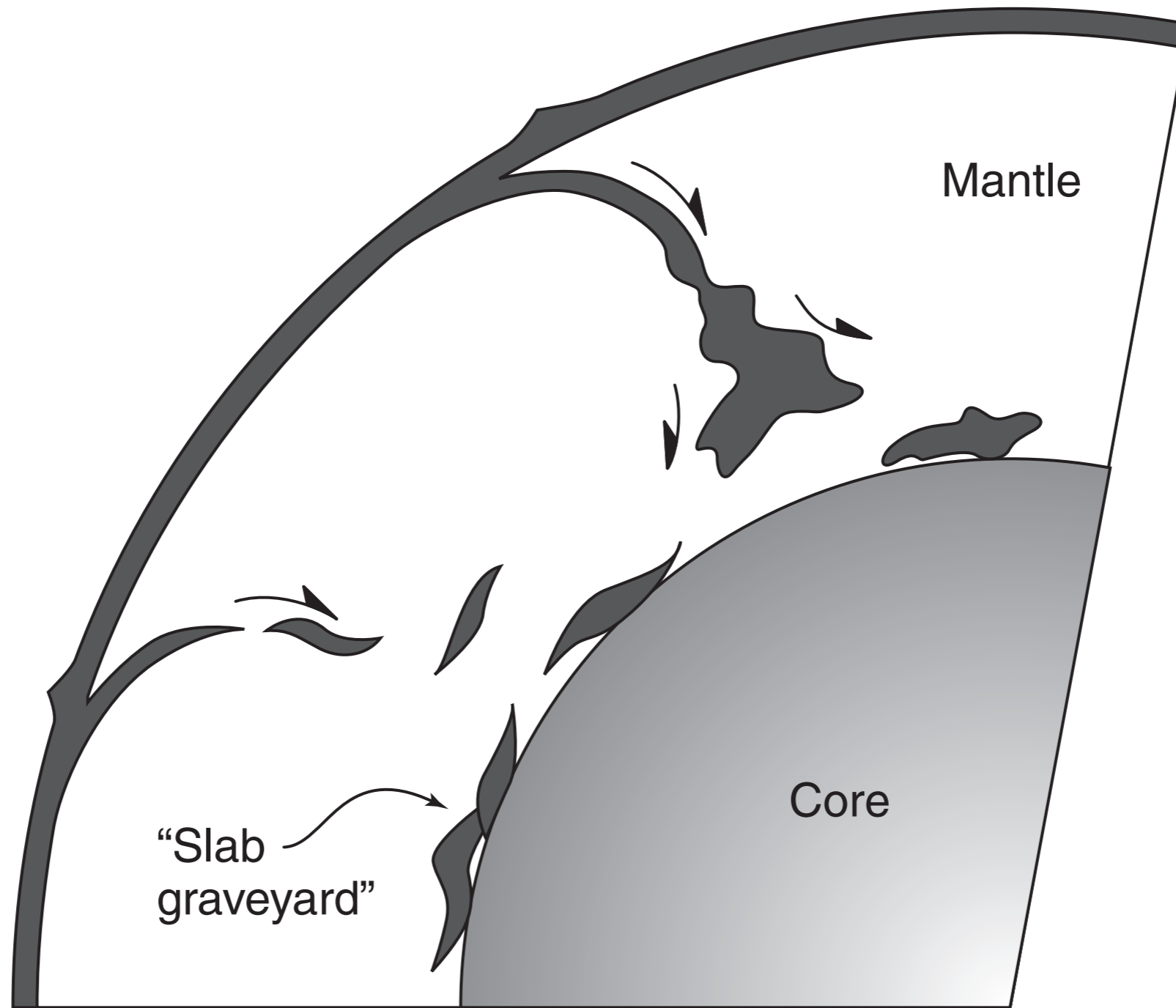
Deep structure of the Alpine area: Center



Deep structure of the Alpine area: West

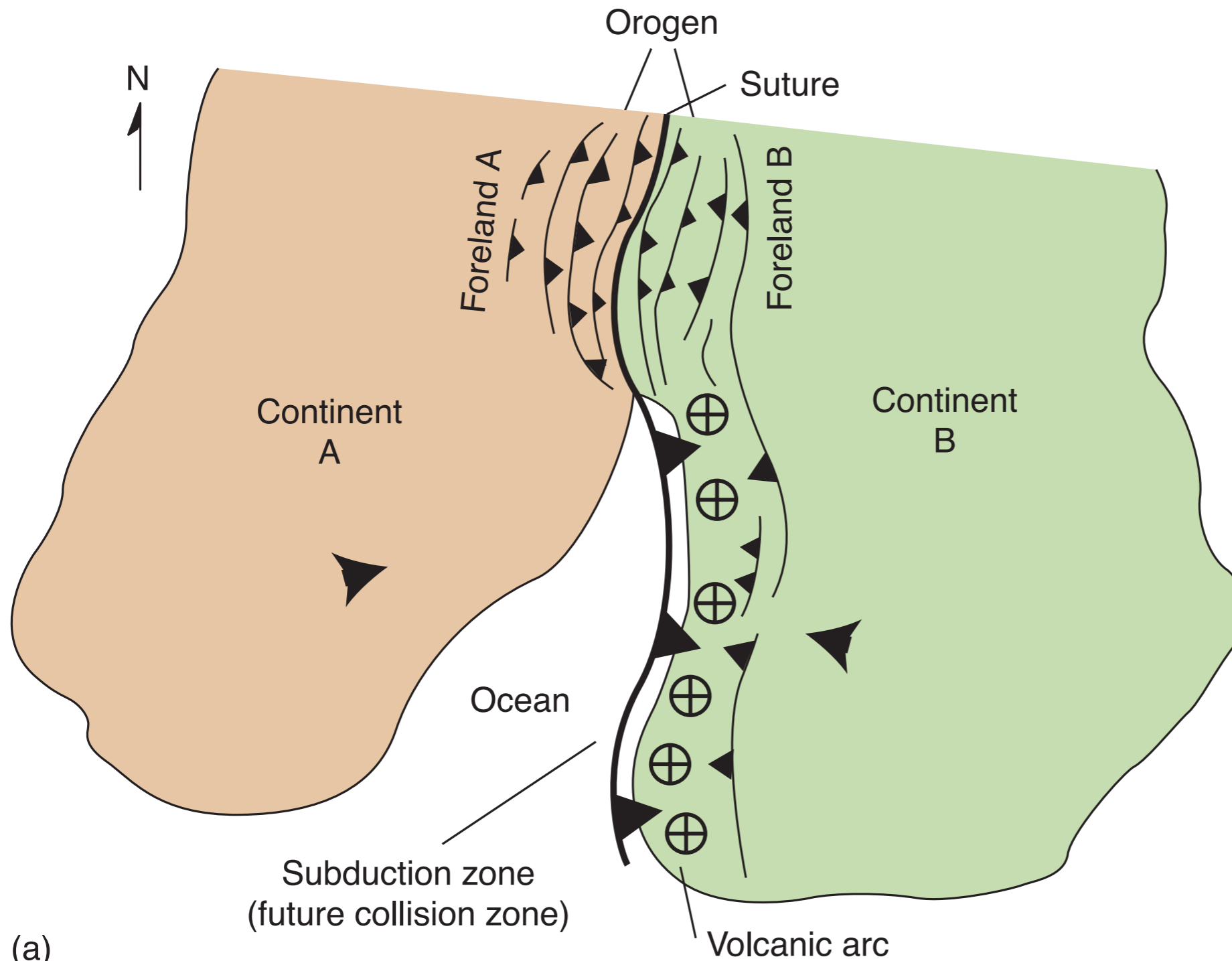


Slab & break-off



Schematic cross section of the Earth illustrating the concept of a slab

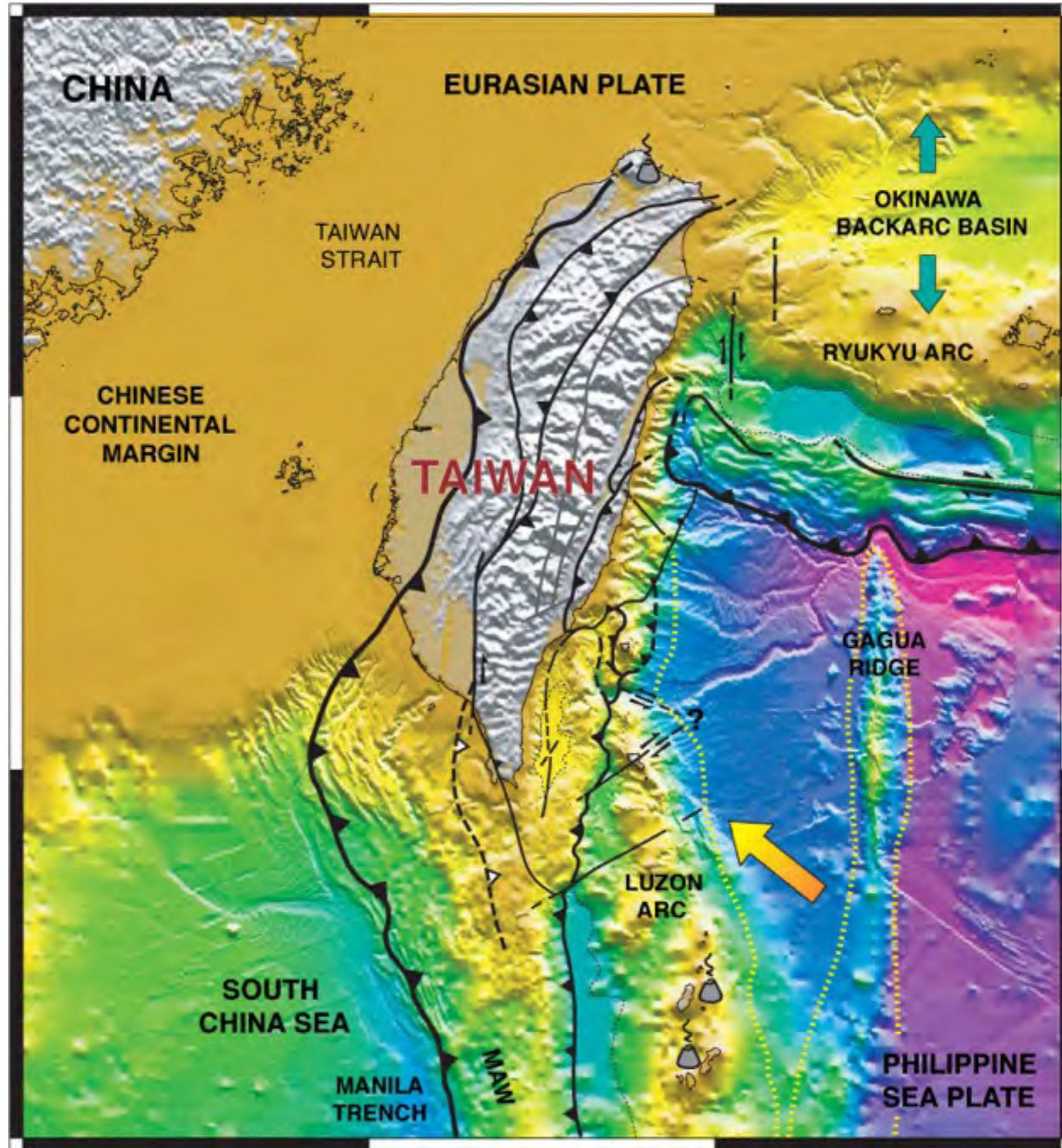
From subduction to collision



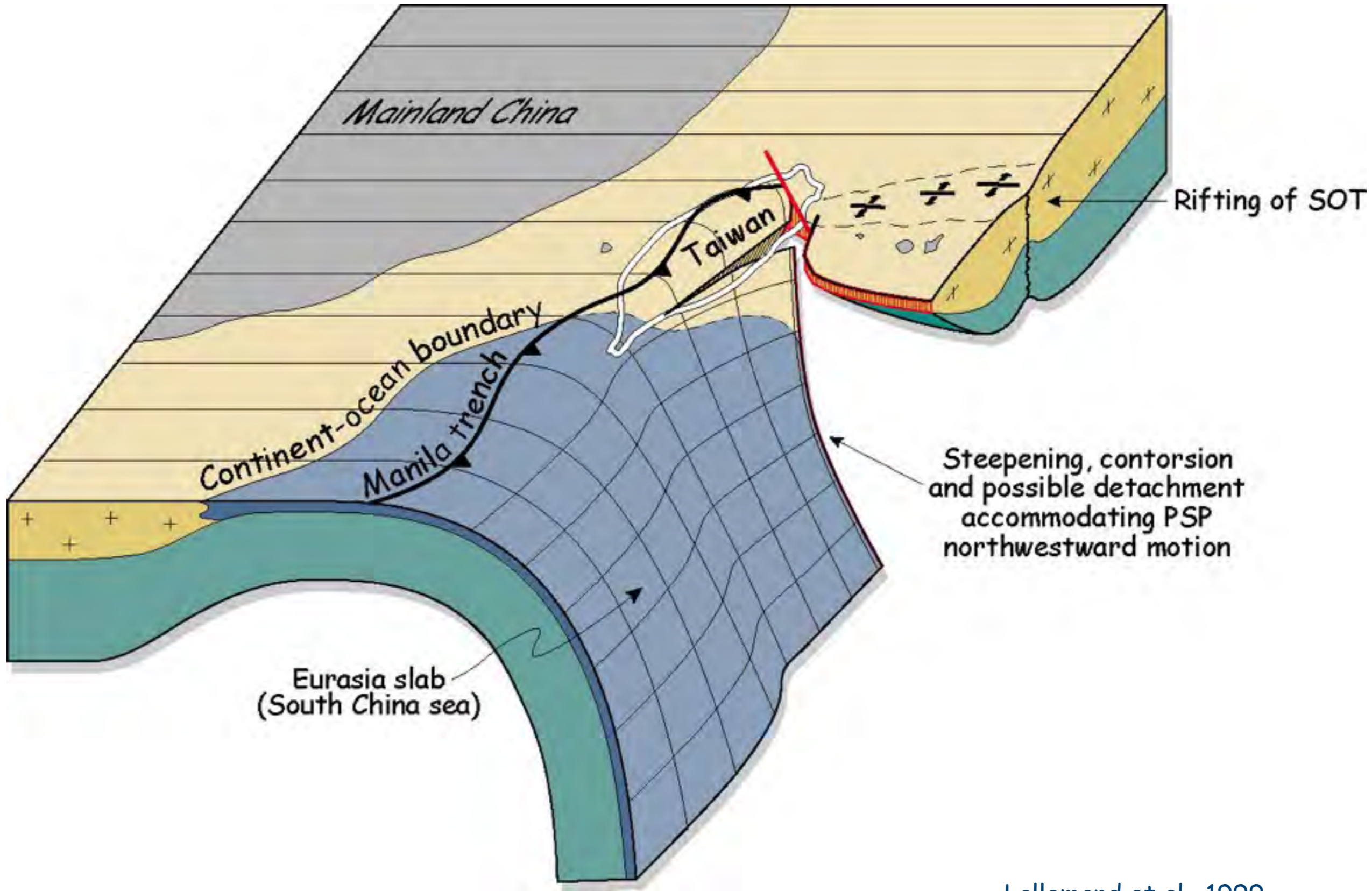
(a)

A map showing a zipper-like collision between two continents. Here, the ocean between the two continents is closing progressively from north to south. In the collision zone, the boundary between what had originally been two separate continents.

From subduction to collision: Taiwan



From subduction to collision: Taiwan



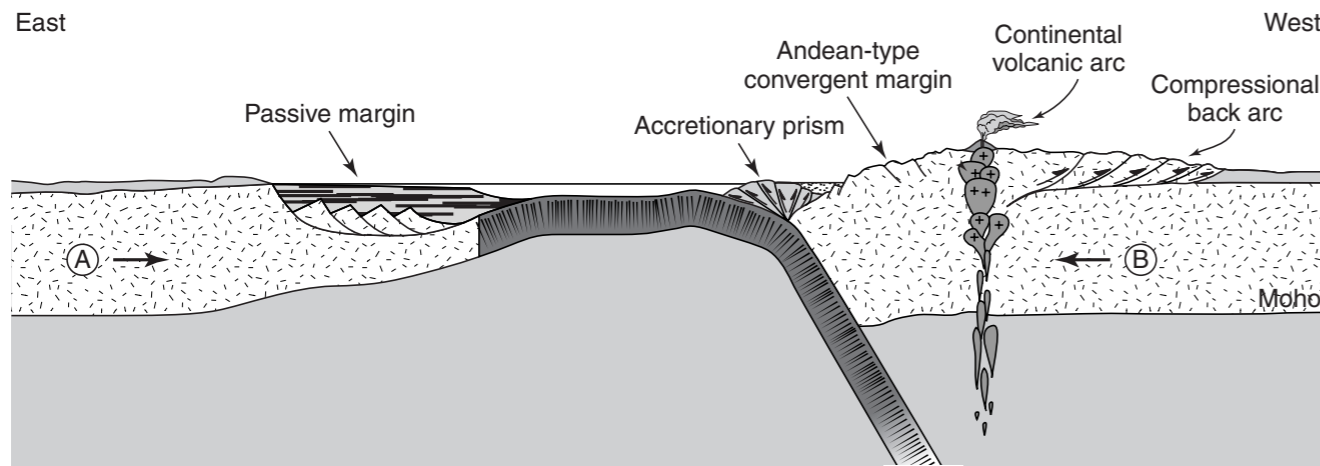
From subduction to collision

Stages in an idealized subduction-collision transition

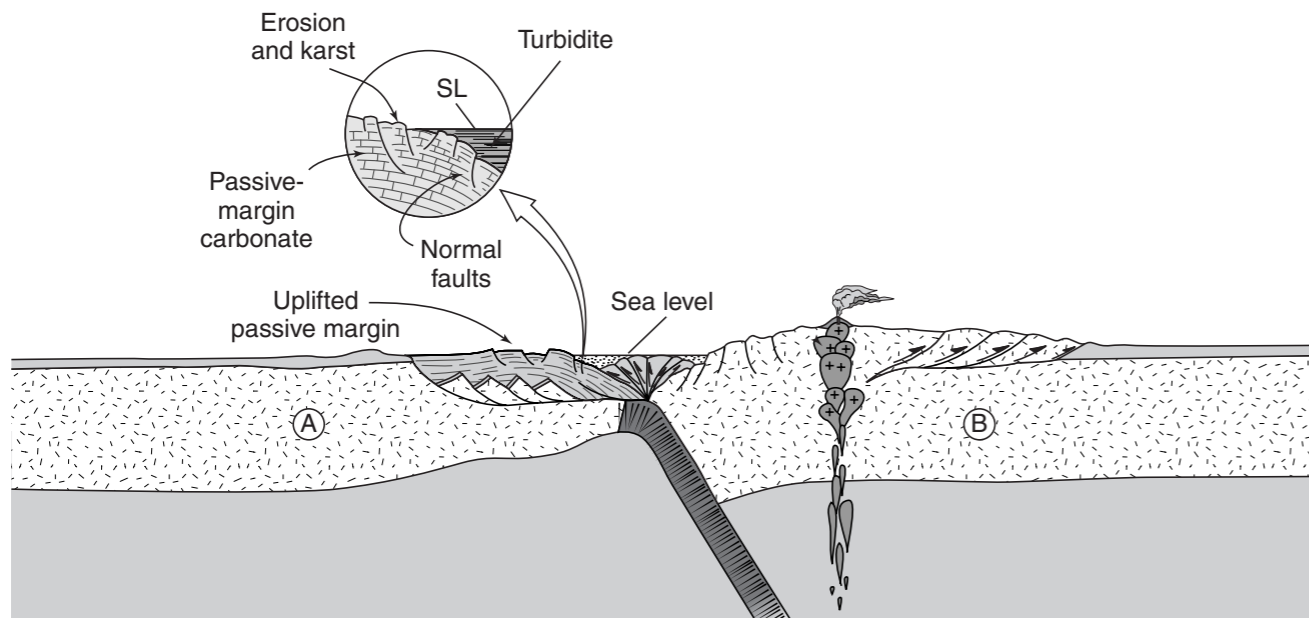
(a) Precollision configuration. Continent A has a passive-margin basin on its east coast, while Continent B has a convergent margin on its west coast. There is a **subduction** of the ocean

(b) During the **continental subduction**, the passive margin is uplifted, and an unconformity (locally, with karst) develops. Turbidites derived from Continent B soon bury this unconformity (see inset). Normal faults break up the strata of the passive-margin basin, due to stretching. But soon, thrusts begin to develop, transporting the deeper parts of the basin over the shallower parts.

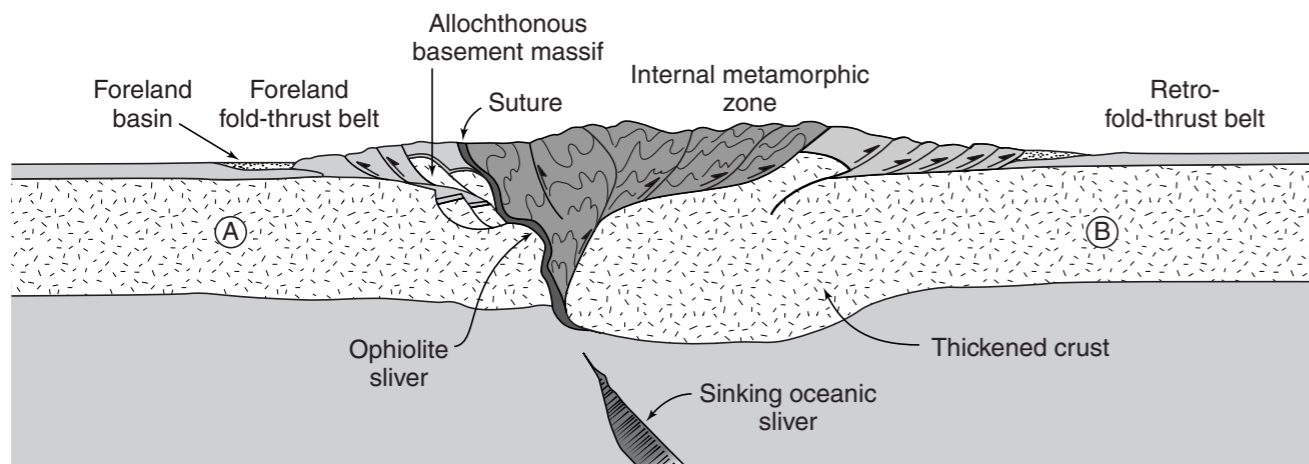
(c) In a **collision orogen**, the subducting **slab has broken off**, a suture has formed, and metamorphic rocks are uplifted and exhumed in the interior of the orogen.



(a)



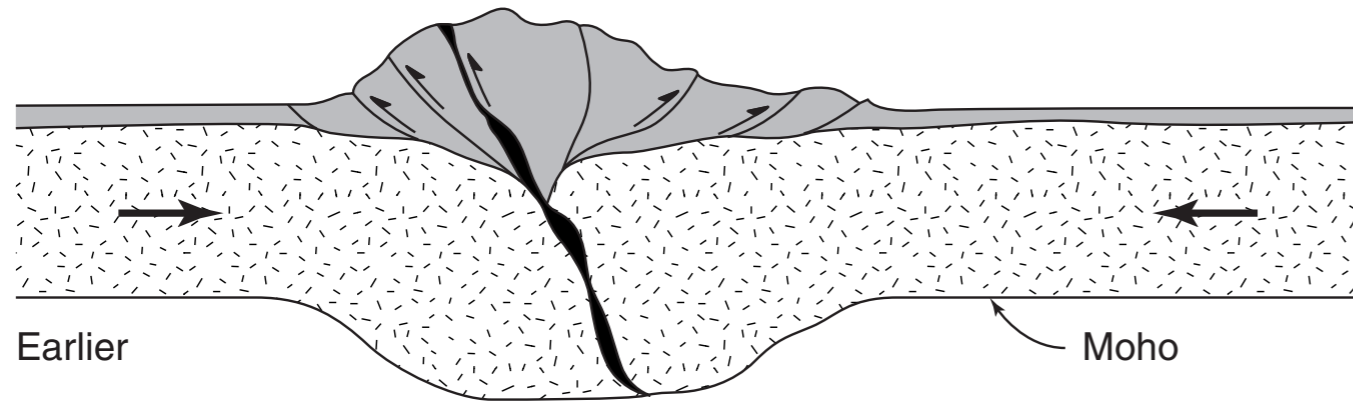
(b)



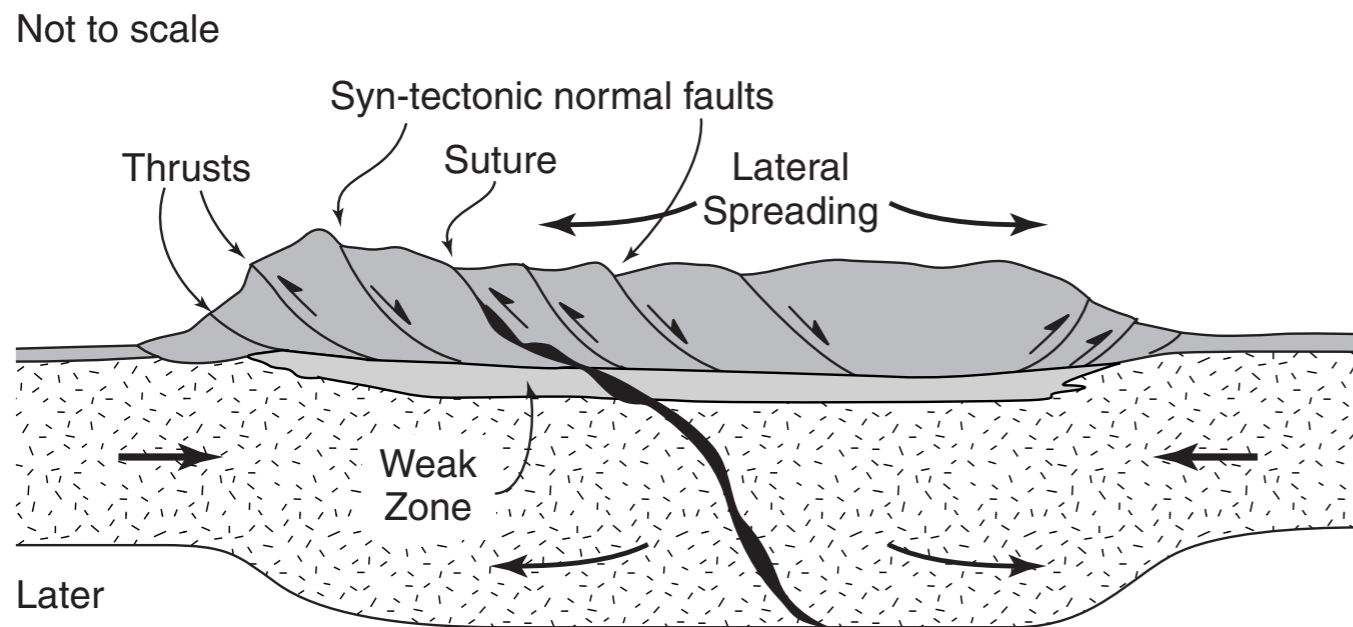
(c)

Collision and post-orogenic processes

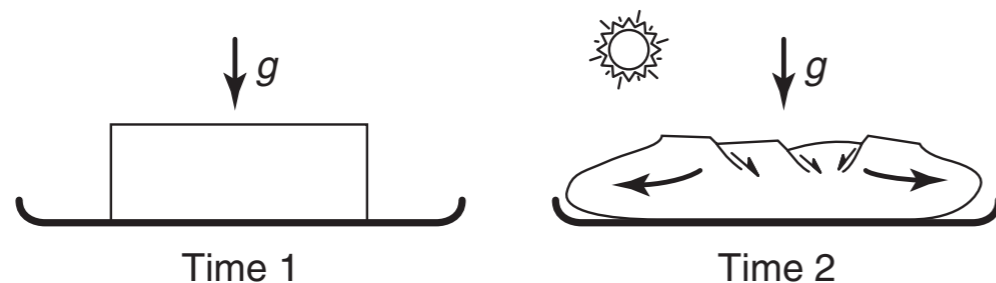
The concept of orogenic collapse



(a) A schematic cross section shows that during an early stage in a collision, the crust thickens by thrusting.



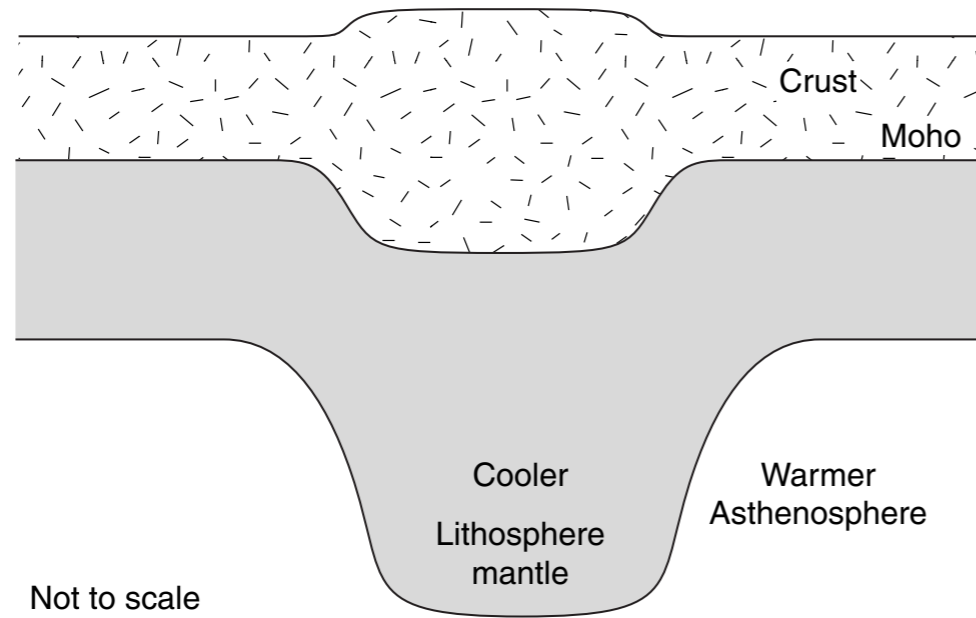
(b) Later, as collapse occurs, extensional faults develop in the upper crust, while plastic flow occurs at depth. This process may contribute to development of a broad plateau.



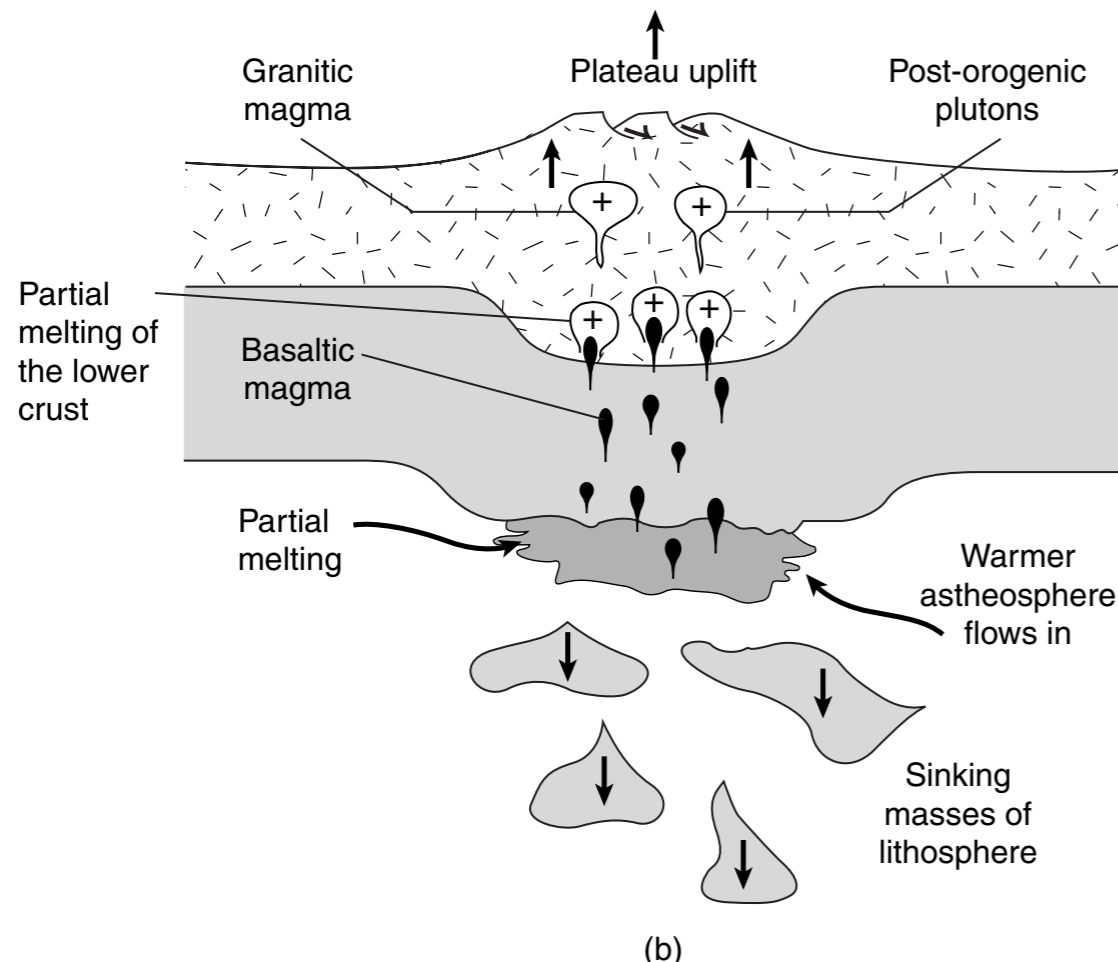
(c) The soft-cheese analogy for extensional collapse. A block of cold cheese can maintain its thickness. If the cheese warms up in the sun, it loses strength and spreads laterally. The rind of the cheese ruptures, and small faults develop

Collision and post-orogenic processes

Post-orogenic plutons and delamination

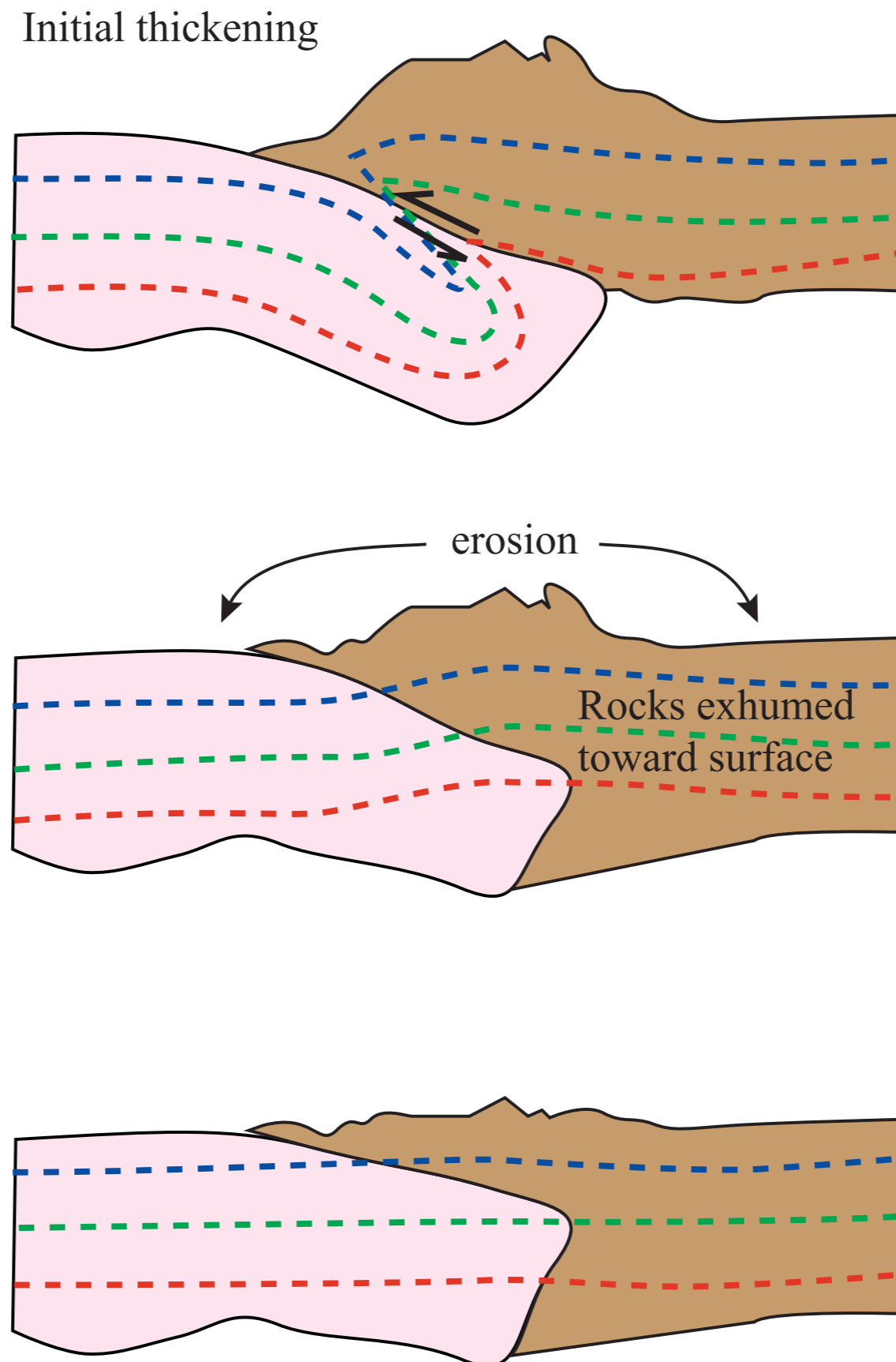


(a) Thickening of lithosphere forms a keel-shaped mass of cool lithosphere to protrude down into the asthenosphere.



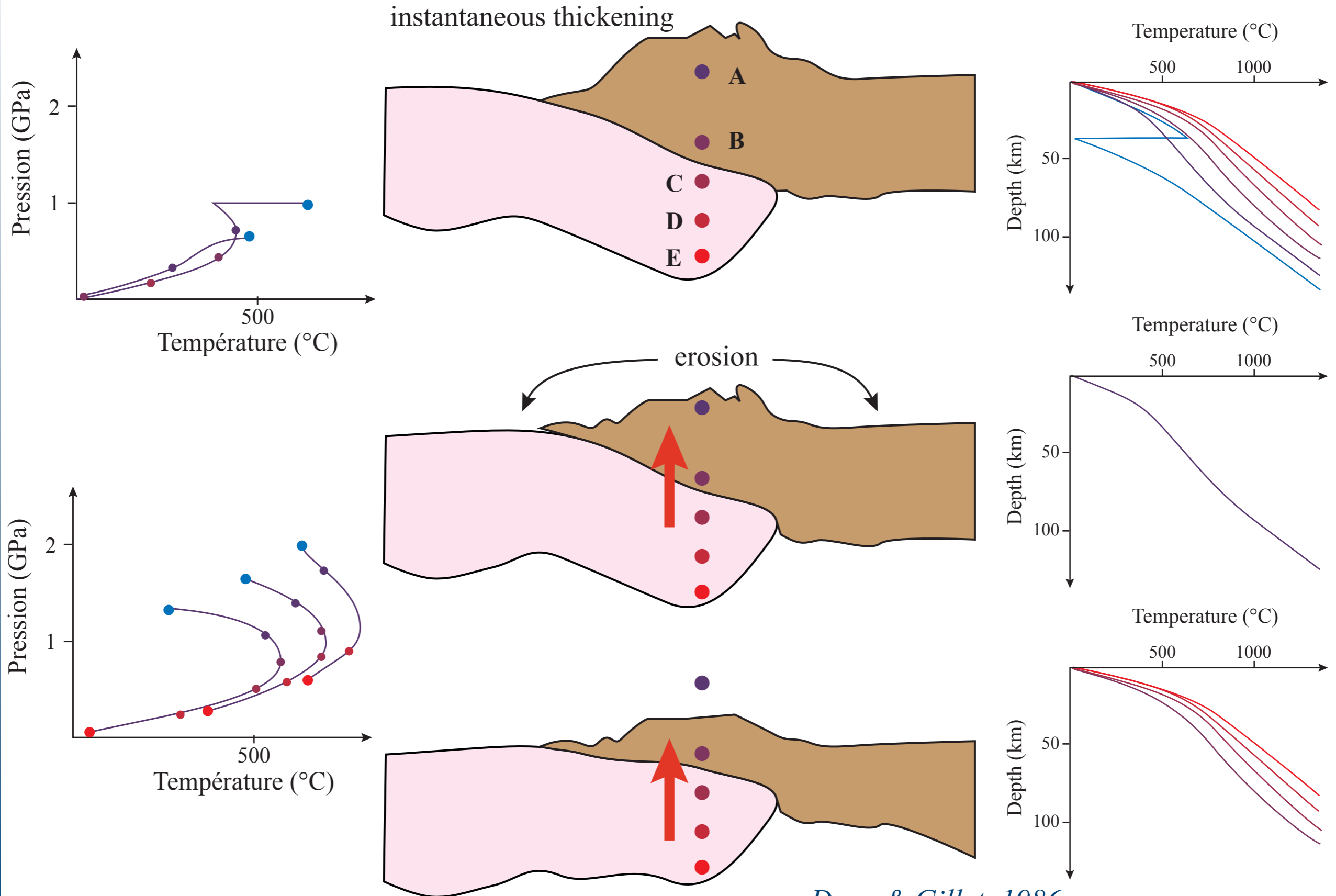
(b) The keel drops off and is replaced by warm asthenosphere, causing partial melting and formation of anorogenic (postorogenic) plutons. The surface of the crust may rise as a consequence

Collision: Thermal Models (1D)



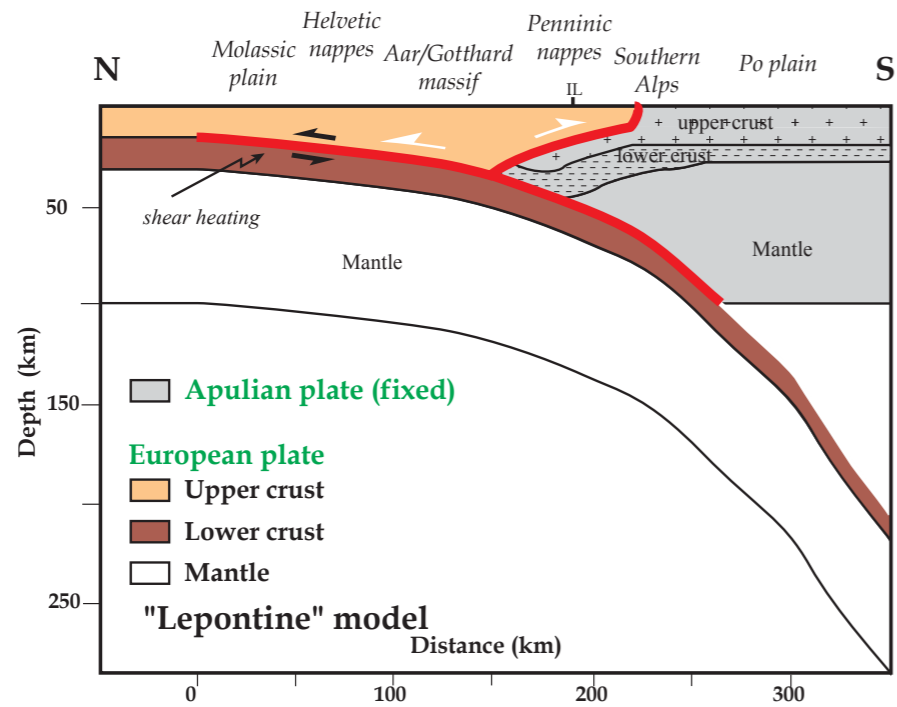
Thompson & England, 1984

Collision: Thermal Models (1D)

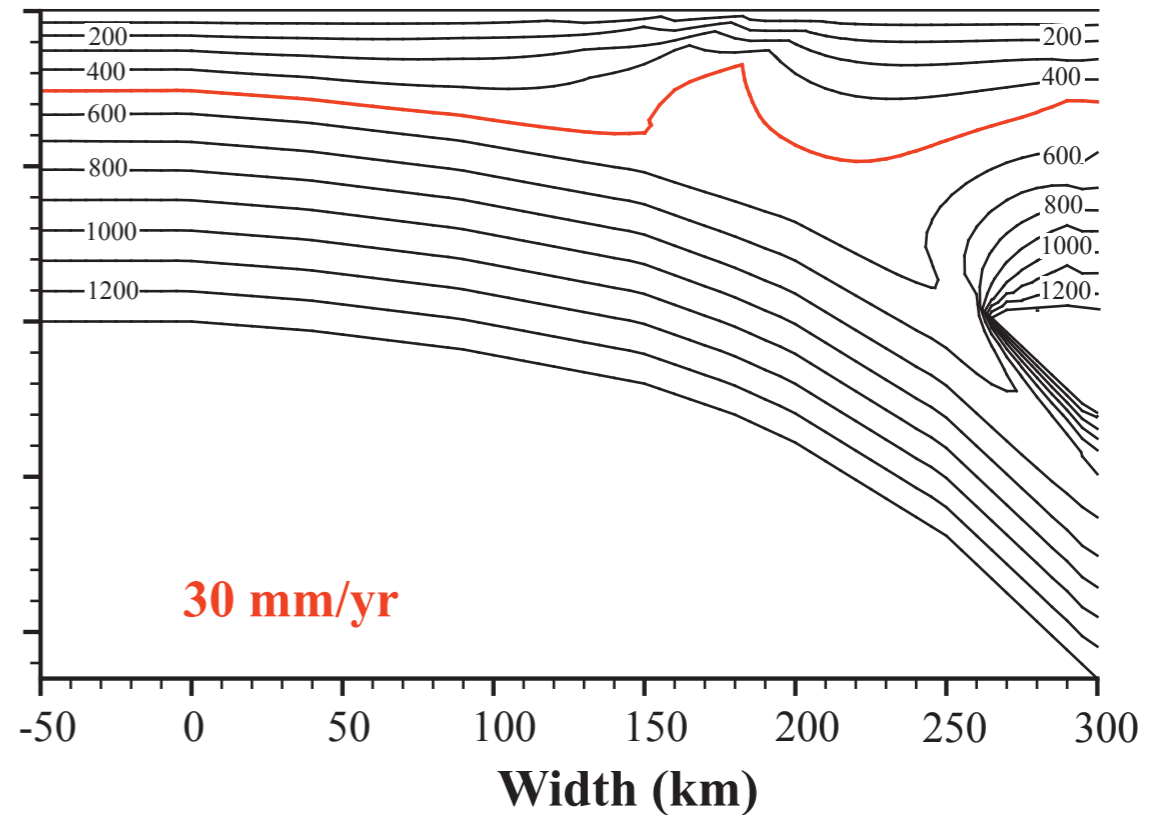
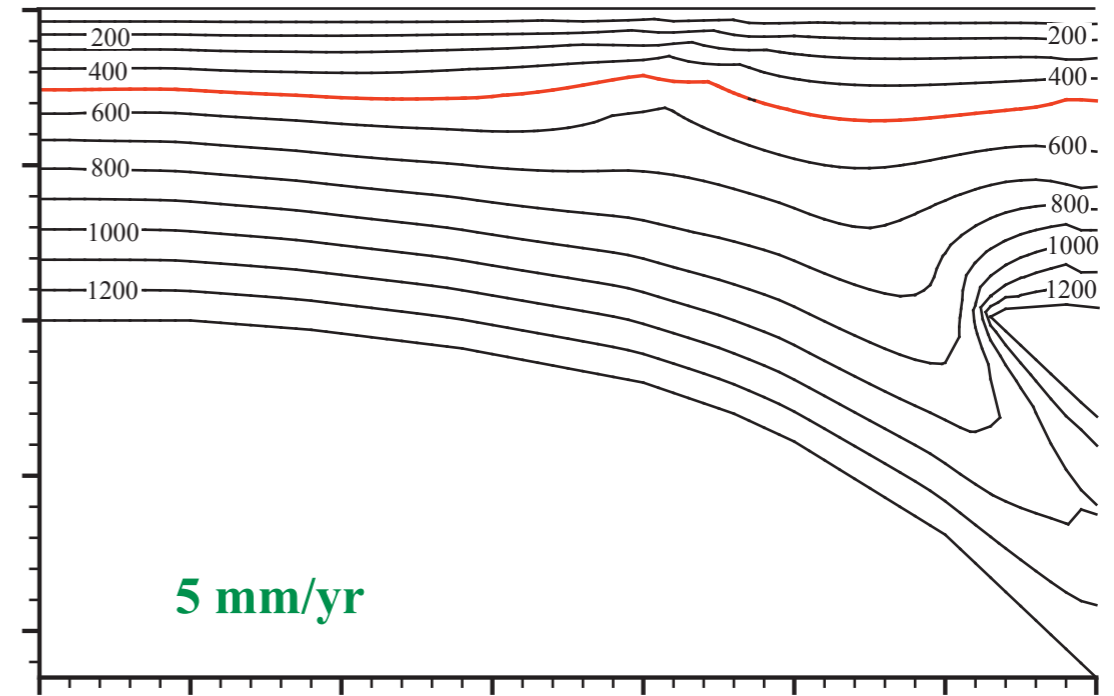


Davy & Gillet, 1986

Collision: Thermal Models (2D)



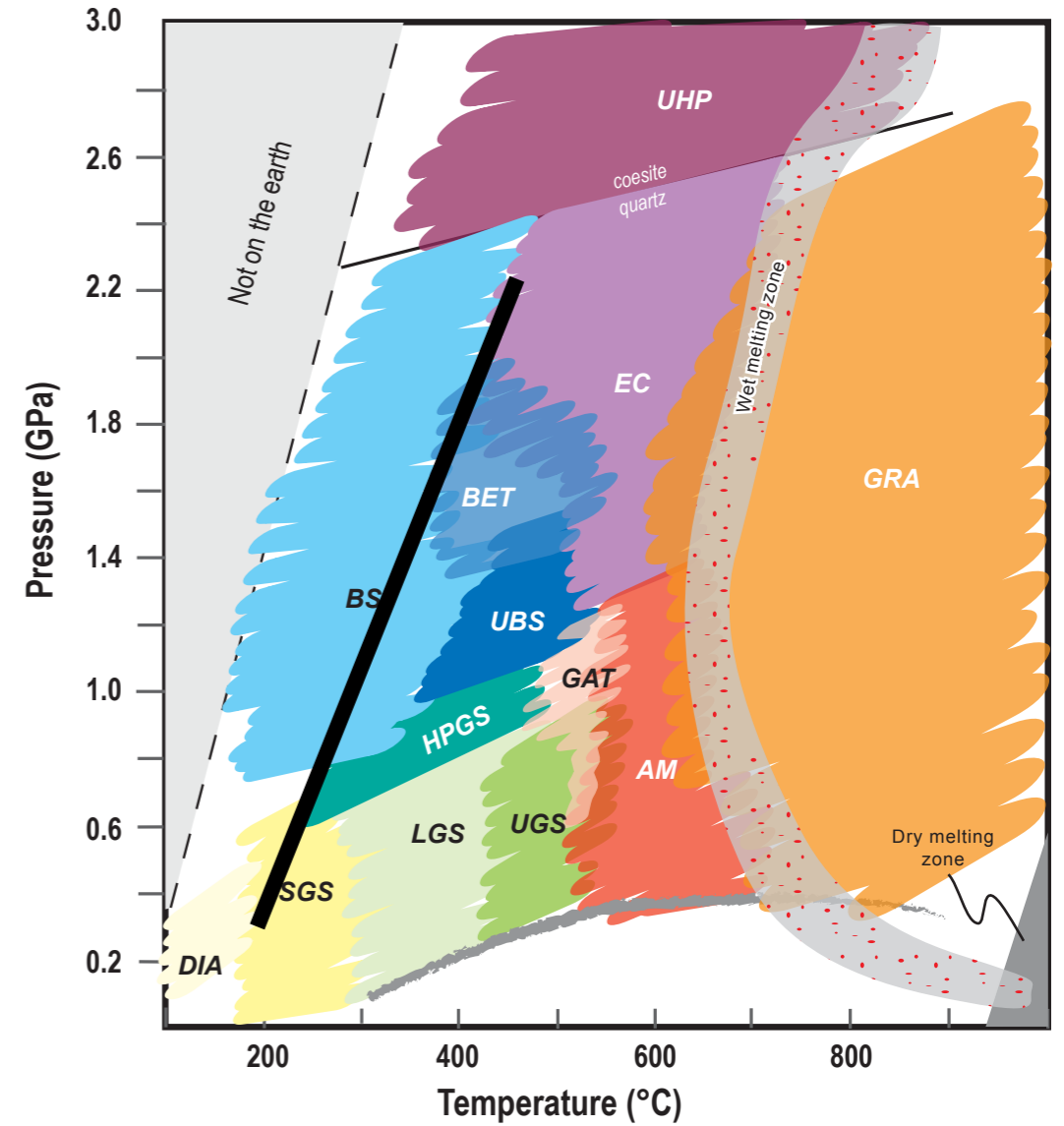
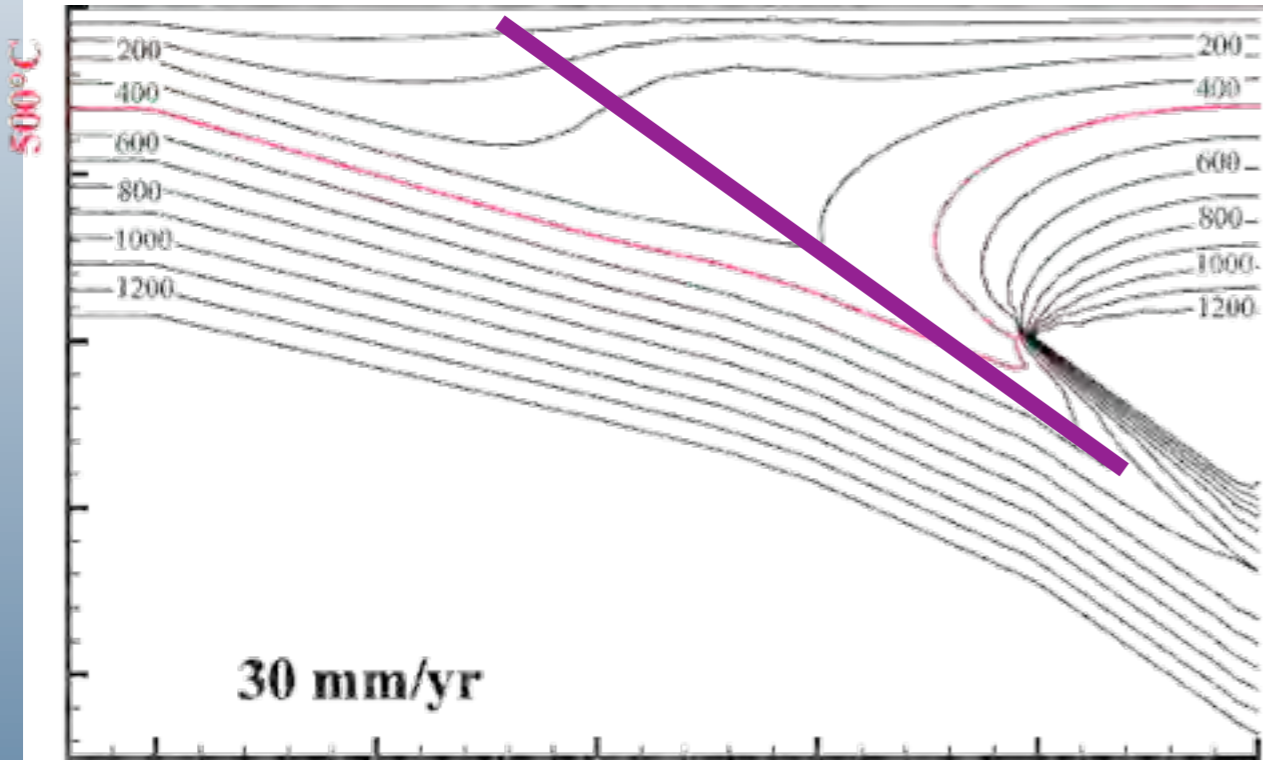
Collision (Alpine-type)



Thermal Models: Subduction vs collision

Subduction

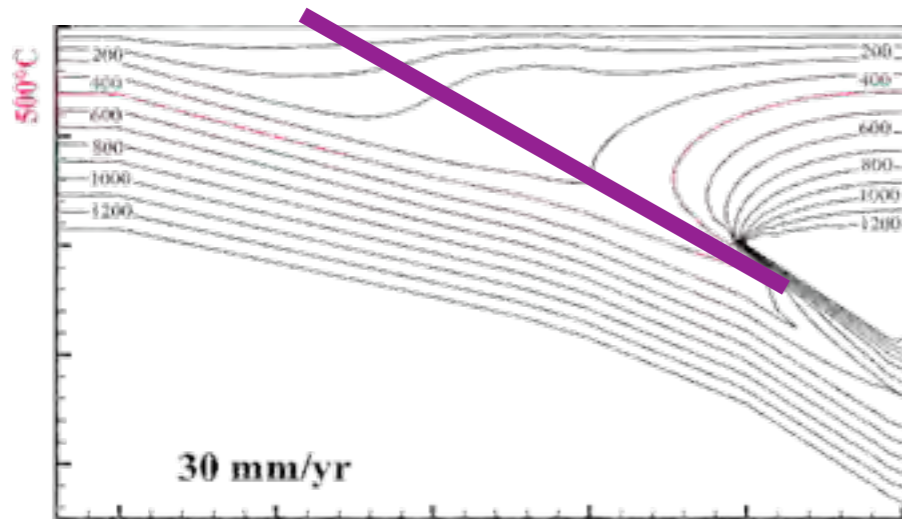
Very "Cold" Gradient
HP-LT $\sim 6^\circ\text{C}/\text{km}$



Thermal Models: Subduction vs collision

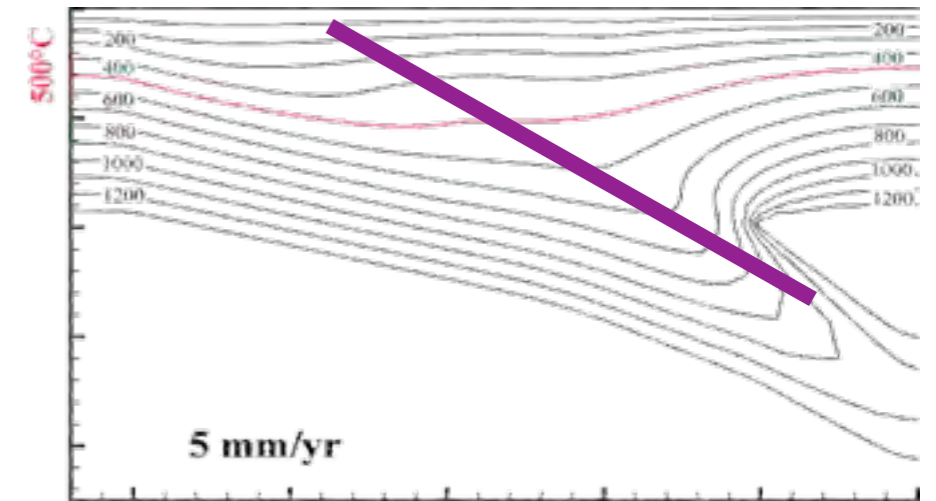
Fast subduction

Very "Cold" Gradient HP-LT ~ 6°C/km

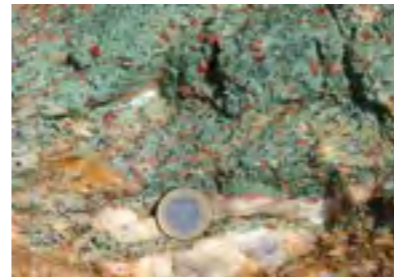


Slow subduction

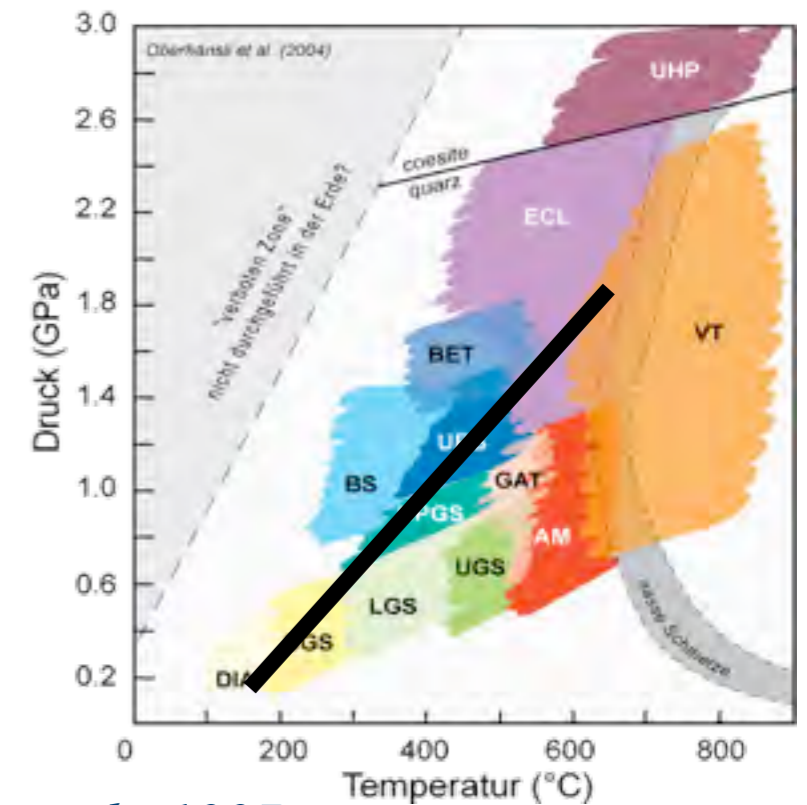
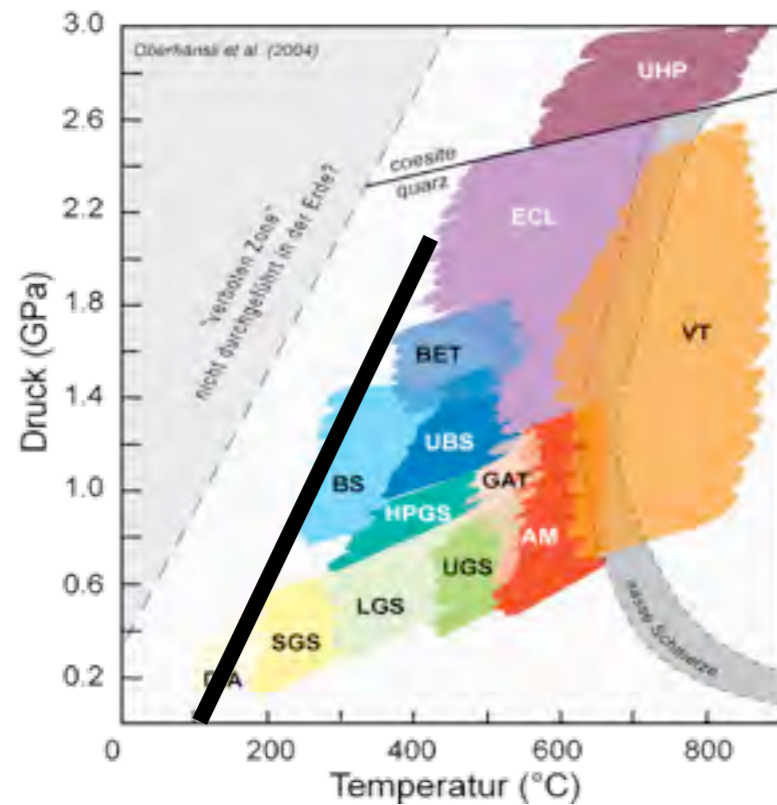
"Cold" Gradient HP-LT ~ 15°C/km



Eclogite



Blueschist

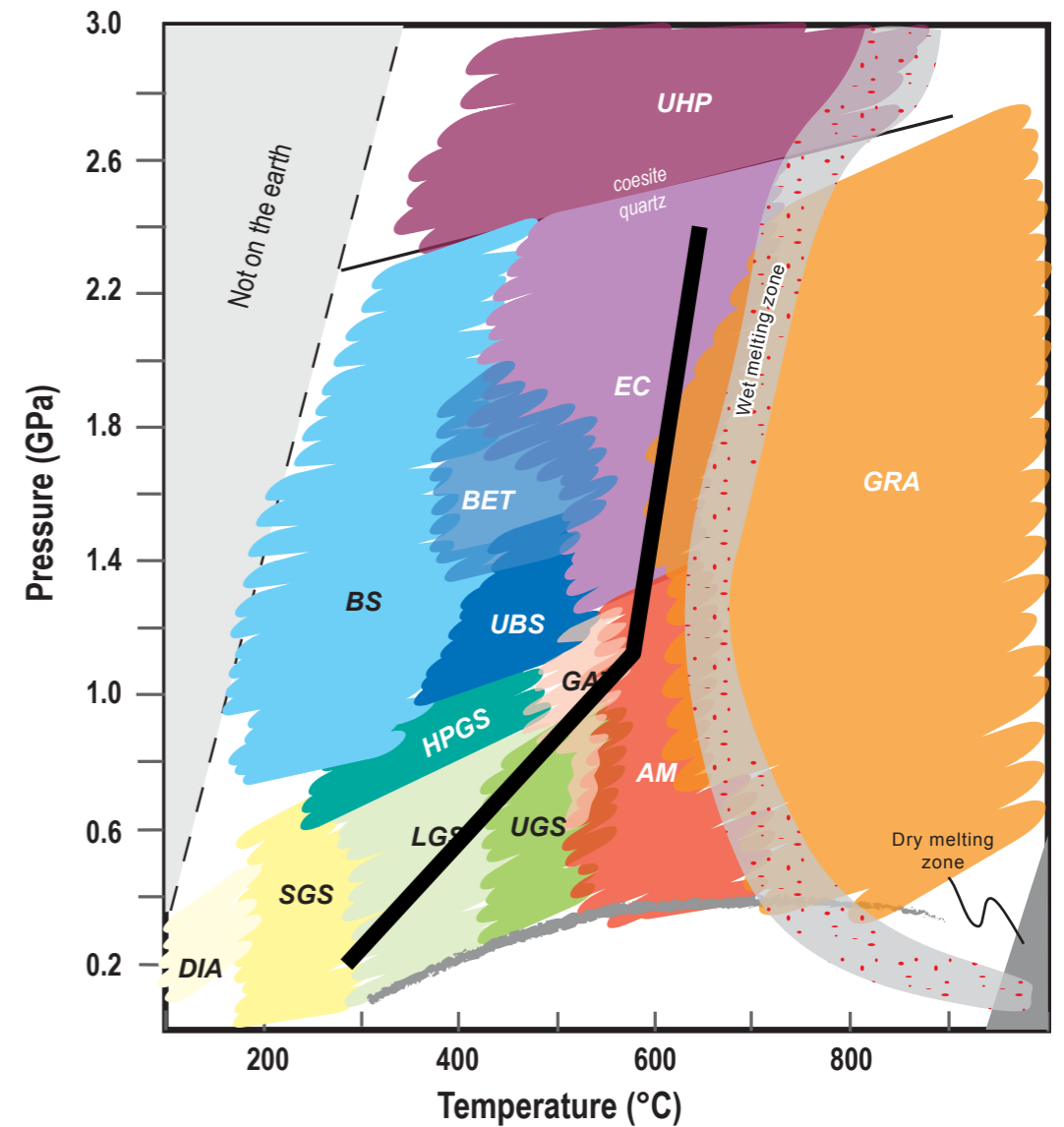
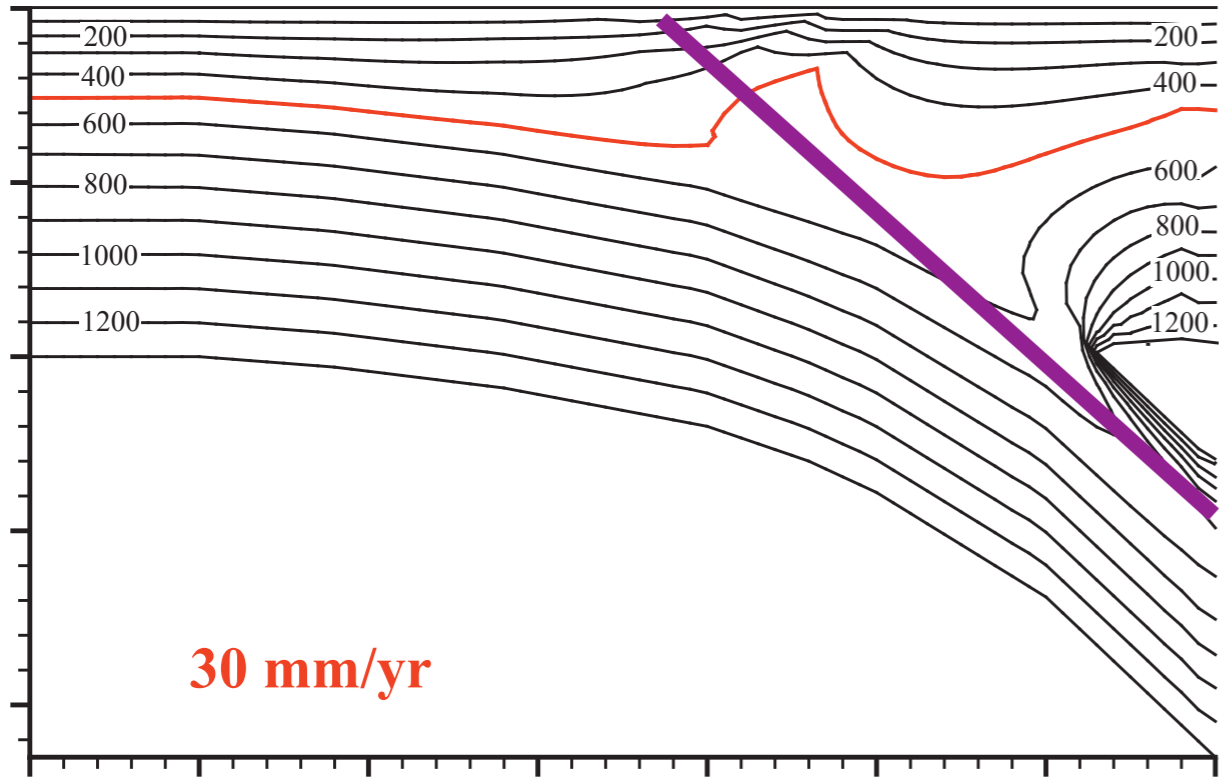


Bousquet et al., 1997

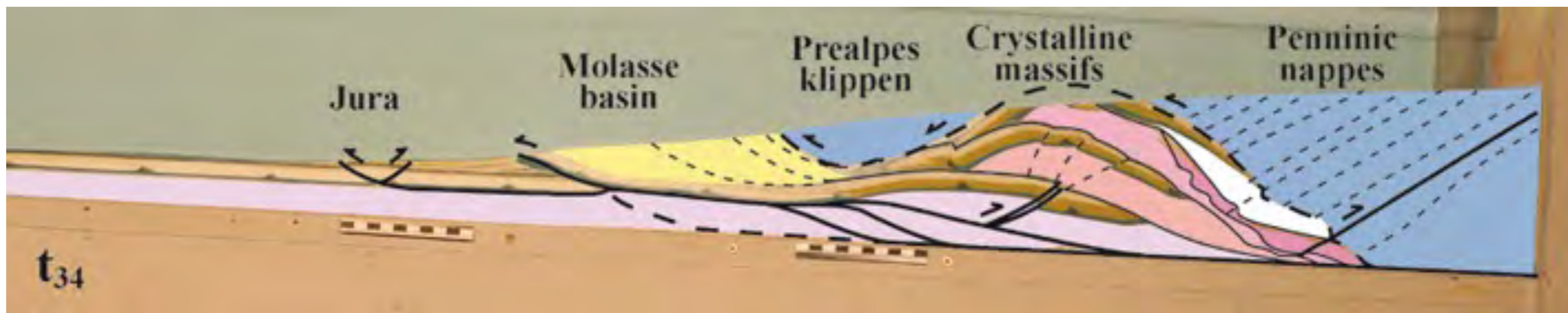
Thermal Models: Subduction vs collision

Continental Subduction

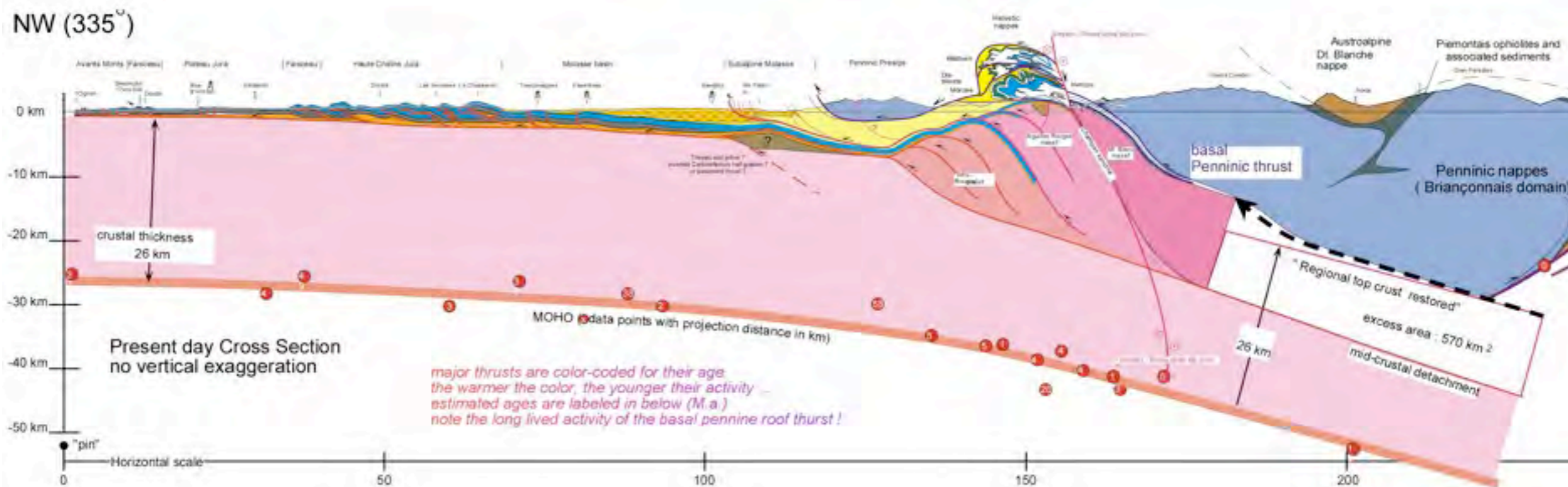
“Middle“ Gradient
HP-MT ~ 15°C/km



Example of continental accretion: the Alps

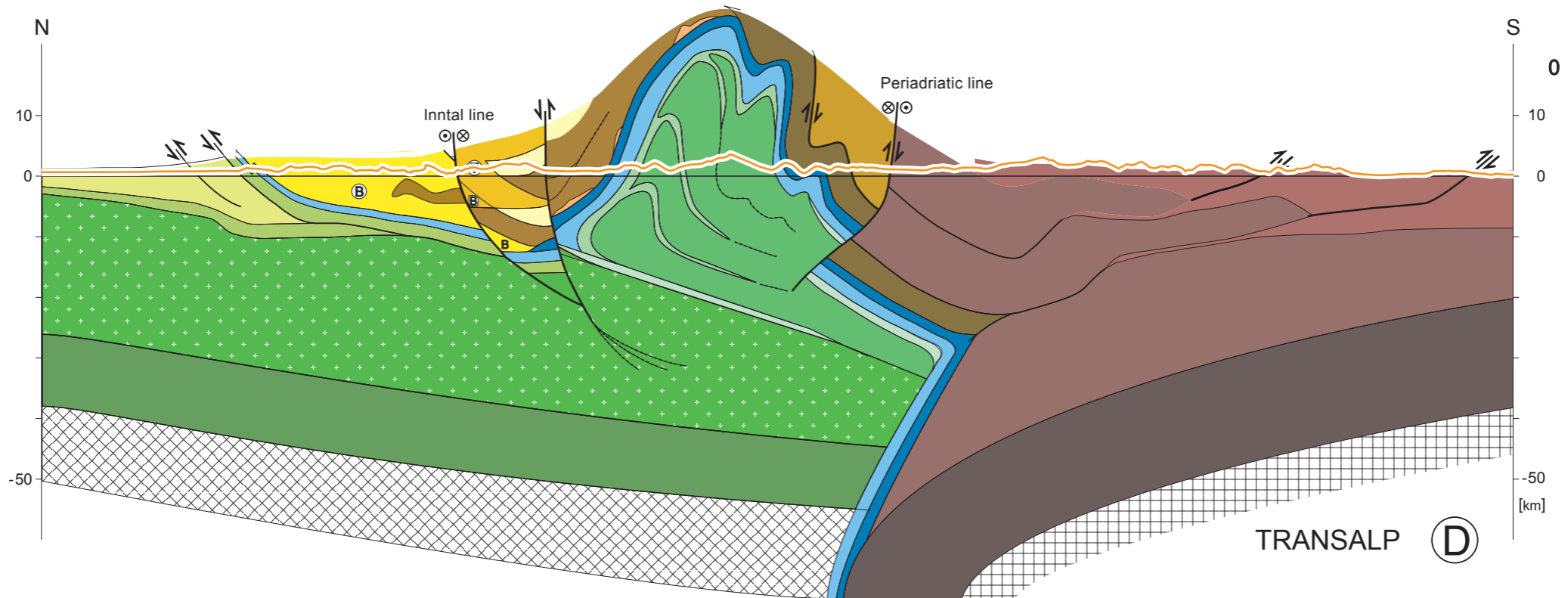
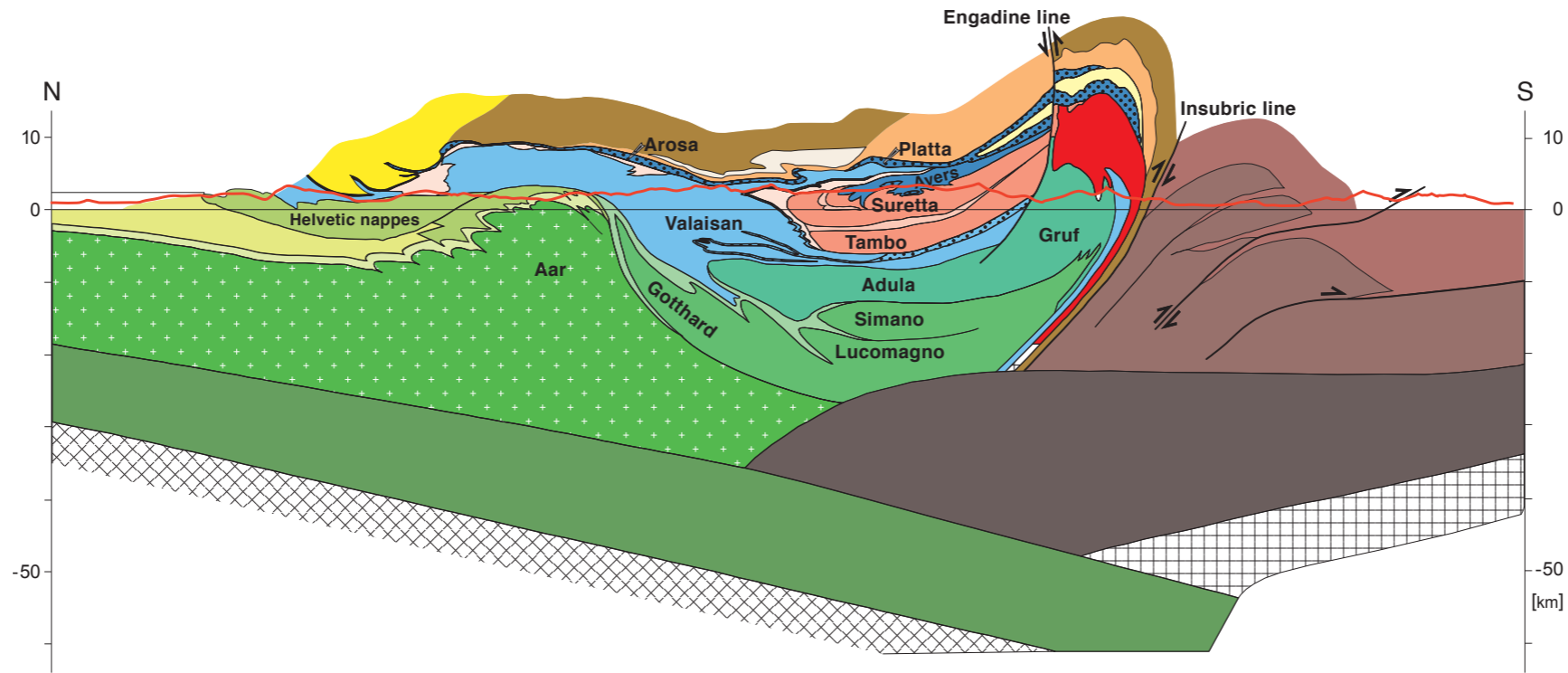


Bonnet et al., 2006



Example of continental accretion: the Alps

Profiles



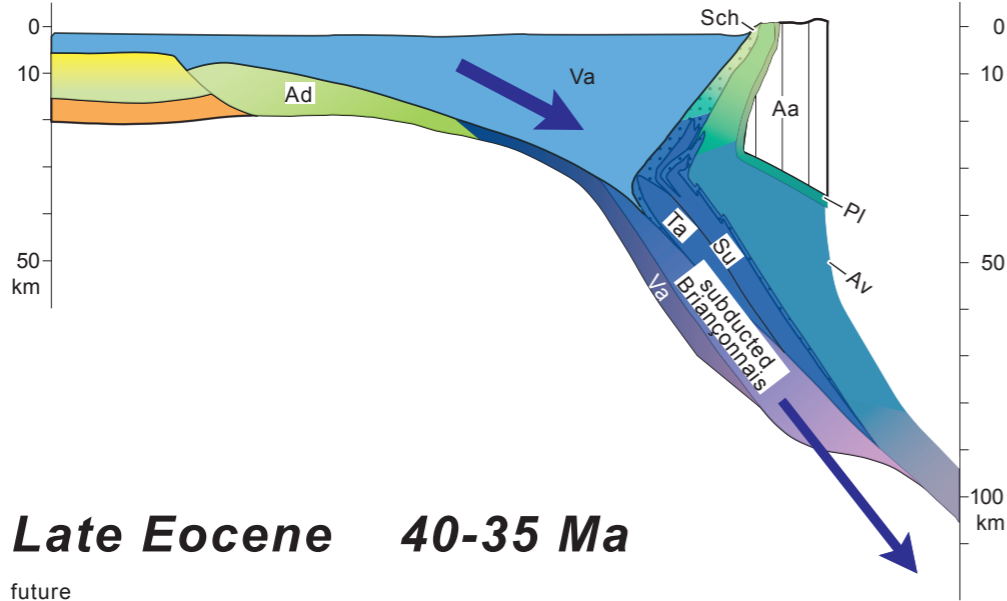
Schmid et al., 2004

Suduction-collision in the Alps

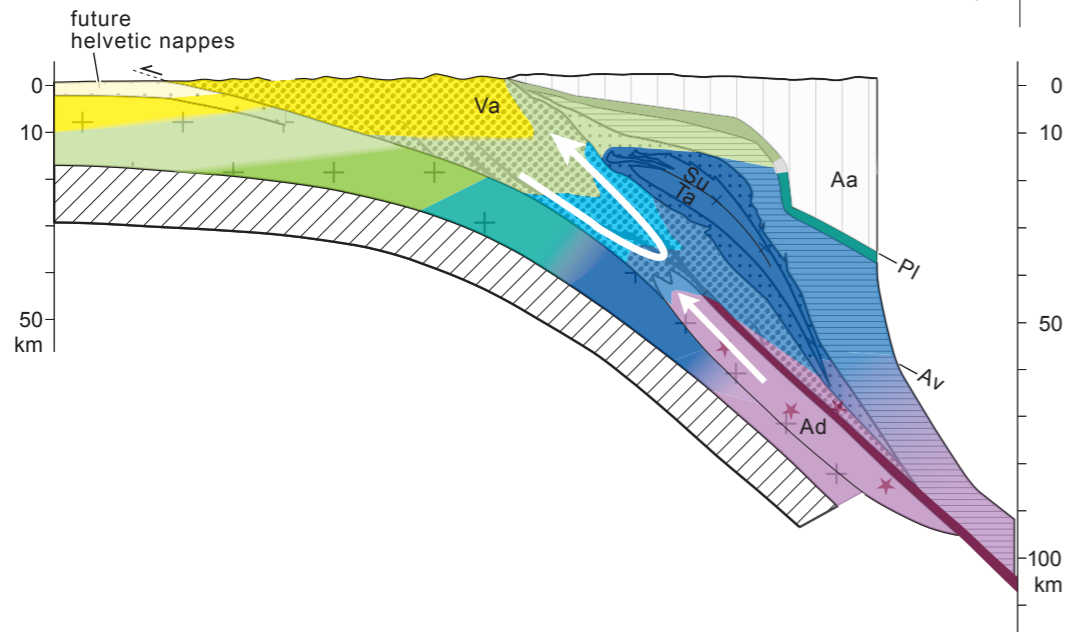
Low temperature regime

High temperature regime

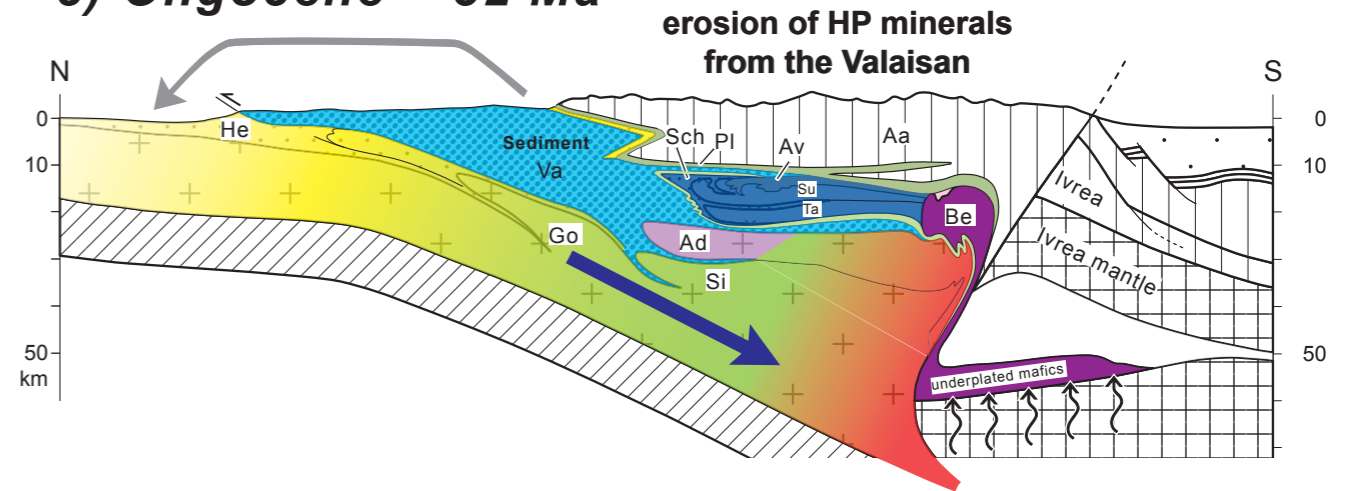
a) Early Eocene 50 Ma



b) Late Eocene 40-35 Ma



c) Oligocene 32 Ma



d) Early Miocene 22 Ma

